Thick & Thin Cloud Top Height Retrieval Algorithm with the Infrared Camera & LIDAR of the JEM-EUSO Space Mission.

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Why it is crucial to detect clouds @ JEM-EUSO FoV ???

- Cloud effect on the EAS detection
- JEM-EUSO cloud efficiency
- 3D clouds UV simulation
- Cloud top height reconstruction



EAS photon interaction

When photons from the Extensive Air Shower interact with the atmosphere, they can be:

- Scattered (reflected)
- Absorbed
- ^D Transmitted



-δscat(dl)

EAS photon interaction

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$$\gamma$$
 N $+E$

$$A \approx N_0 e^{-\delta abs(dl)}$$

These processes strongly depend on the presence of clouds, leading to systematic errors in the reconstruction of the UHECRs and EECRs.

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JEM-EUSO traverses over various atmospheric situations within its huge observation area. Space-based telescopes can observe EAS under certain cloudy conditions.



JEM-EUSO AMS payloads: I. Infrared Camera responsability of Spain II. LIDAR (LIght Detection And Ranging) responsibility of Japan & Switzerland



Distribution of the shower maximum depending on the incident angle



In energy range of interest the depth of EAS development (Xmax) Only varies by ~100 g/cm2. Hmax is dominantly determined by the zenith angle.

Stratus



Cirrus



Thin clouds effects in the UHECRs energy spectrum



Detected EAS photons under different cloudy conditions



Sáez Cano et al., Advances in Space Reseach, 2014.

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EAS FS image affected by clouds JEM-EUSO coll, Exp. Astronomy 2014, in press



Cloud impact (TOVS). The JEM-EUSO Coll. ApP, 2013.



Clear atmosphere

~ clear atmosphere

Observable with reduced signals

Observable for high ZA

Hardly observable

Cloud efficiency for E> 10^{19.5} eV

Cloud-top I	Optical depth $\tau_{\rm C}$			
Altitude	0.05	0.5	1.5	5
$H_{\rm C} = 10 \ \rm km$	90%	70%	26%	18%
$H_{\rm C} = 7.5 \ \rm km$	89%	74%	43%	37%
$H_{\rm C} = 5 \ \rm km$	89%	82%	69%	66%
$H_{\rm C} = 2.5 \ \rm km$	90%	88%	89%	88%

The JEM-EUSO Coll, Astroparticle Physics, 2013



Satellite Data Simulation Unit (SDSU)

SDSU v2:

- → Developed in Nagoya University.
- Simulates from 0.17 1000 μ m.
 - IR Model: Nakajima et al. (2003).

SDSU code inputs provides for each pixel:

height, pressure, air temperature, relative humidity, cloud water content, rain water content, cloud ice content, snow ice content, graupel ice content, and hail ice content.



Thick and thin cloud 3D UV simulation

But which is the cloud top altitude?



EAS fluorescence photons progations in a 3D cloud EAS of 60°







Lidar OD vs layer Lidar Signal vs layer Lidar Nghe Lidar OD 400 Legend Legend WC=1 [mg/m^3] WC=1 [mg/m^3] WC=5 [mg/m^3] WC=5 [mg/m^3] WC=10 [mg/m^3] 10-1 WC=10 [mg/m^3] 300 WC=20 [ma/m^3] WC=20 [ma/m^3] 200 100 10-2 1040 1042 1044 1046 1048 1050 1052 1054 1050 1054 1040 1042 1044 1046 1048 1052 Height [GTU] Height [GTU]

J.A. MORALES DE LOS RÍOS, JEM-EUSO INTERNATIONAL MEETING, TENERIFE 2013



Reconstructed Height



Thin Ice-Cloud Optical Depth Study.

$$\beta = \beta_{ground} \times (1 - \tau) + \beta_{cloud} \times (\tau)$$

$\tau = \tau abs + \tau sca$

OD calculates using parametrization from Qiang Fu and K.N. Liou, in the paper titlet Parametrization of the Radiative Properties of Cirrus Clouds.

Where the Optical depht for scattering is defined as;

$$\tau sca = \beta \times L$$

and β is a parameterization like,

$$\beta = \sum_{n=0}^{N} a_n / De^n \qquad (6.9)$$

in witch De is the mean effective radius defined for spherical water droplets as follows,

$$De = \frac{\int_{Lmin}^{Lmax} D \times DL(L) \times n(L) dL}{\int_{Lmin}^{Lmax} DL(L) \times n(L) dL}$$
(6.10)

being D the with of an ice crystal, n(L) denotes the ice crystal size distribution, and Lmin and Lmax are the minimum and maximum lengths of ice crystals, respectively. The geometric cross section area for randomly oriented hexagonal ice crystals generally deviates from DL(L), and may be expressed by (Takano and Liou 1989).

$$DL = \frac{3}{2}D(\frac{\sqrt{3}}{4}D + L)$$

TABLE 3. Values of empirical coefficients in Eq. (2.9) for the parameterization of the extinction coefficient β (m⁻¹). Note that the coefficients are determined by using an ice density ρ_i of 0.9167 g cm⁻³. For other ice densities ρ_i^* , they should be adjusted by a factor ρ_i/ρ_i^* . (The units for D_e and *IWC* are in μ m and g m⁻³, respectively.)

Band <i>i</i>	<i>a</i> ₀	<i>a</i> ₁	a	
16	-6.656×10^{-3}	3.686	0.0	
7	-7.770×10^{-3}	3.734	11.85	
8	$-8.088 imes 10^{-3}$	3.717	17.17	
9	-8.441×10^{-3}	3.715	19.48	
10	-9.061×10^{-3}	3.741	26.48	
11	-9.609×10^{-3}	3.768	34.11	
12	-1.153×10^{-2}	4.109	17.32	
13	$-8.294 imes 10^{-3}$	3.925	1.315	
14	-1.026×10^{-2}	4.105	16.36	
15	-1.151×10^{-2}	4.182	31.13	
16	$-1.704 imes 10^{-2}$	4.830	16.27	
17	-1.741×10^{-2}	5.541	-58.42	
18	-7.752×10^{-3}	4.624	-42.01	

(6.8)

(6.11)

TDiff Clear Sky vs IWC, Lenght=500m Tground=300K Tcloud=235.5K



Cloud Mask Strategy under development



Cloud Mask Real Simulation (Preliminary).

CloudMask 1->Clear Sky, 2-> Thin Cloud, 3->Thick Cloud



Final results to come: Morales de los Ríos PhD Thesis 2014

Summary and conclusions

- UHECRs can be observed in certain cloudy conditions by space based telescopes with a proper AMS (IR camera +LIDAR)
 - Depending on clouds and showers properties
- For JEM-EUSO, an atmospheric monitoring system is needed
- For JEM-EUSO, around 72% of events in cloudy cases can be used for analysis
- A module to simulate 3D clouds in UV has been performed
- A module to propagate EAS photons in 3D clouds is being performed

Summary and conclusions

- To fulfill IR camera requirements, the cloud algorithm has to properly retrieve the height of the clouds in the whole FoV.
 - Overall NETD Budget < 3K (500 m)
- This CLOUD TOP HEIGHT RETRIEVAL ALGORITHM retrieve the cloud top heights of thin and thick clouds without any intermedia steps that innecesarily increases the NETD budget.
- The HEIGHT RETRIEVAL ALGORITHM of JEM-EUSO is based on SDSU simulation and considers a LIDAR shot in the IR camera image.
- Cloud mask to distinghuish between clear sky, thin clouds and thick clouds is also successfully under development

BACKUP SLIDES

• Cloud Top Height Reconstruction • Adiabatic atmosphere aproach, with Lidar signal.

$$\frac{d\rho}{dz} = -\frac{\rho}{g_n}$$

First we have an Hydrostatic equation:
...where rho=density, z= height, gn=Std gravity.

• Integrating...

• If liquiq:
$$p = \rho \times g_n \times z + p_0$$

• If gas:
$$p = p_0 * e^{\frac{-M \times g_n \times z}{R \times t}}$$
 •

$$p = \rho \times \frac{R}{M} \times t$$

Using...

 And then, defining the atmosphere pressure with this adiabatic aproximation:

$$p \times t^{\frac{-\alpha}{(\alpha-1)}} = Ctte$$

- ... where:
- gn=std gravity= 9.80665 m/s^2.
- M=Air molar mass= 0.02897 kg/mol
- R= gast ctte= 8.3144621 J/(mol*k)
- alpha = heat cap ratio=7/5
- P0= pressure at sea level= 101325 Pa.

 Using one reference height position got by the Lidar, and the cloud temperature in the same position, we can calculate the adiabatic constant, and then reconstruct height of other pixels.

Photon propagation module for 3D clouds: How we treat light from EAS

FLUORESCENCE:

Only directly emitted light to the telescope is considered

CHERENKOV:

Only first order of scattering is considered

Cherenkov bunches:

<u>Considerations:</u>

- No emitted photons in the direction of the telescope.
- Rayleigh scattering has been considered.

• <u>How we proceed:</u>

1) We calculate the cell of the atmosphere where the bunch is located, and the number of photons in the bunch

For each Cherenkov sample:

2) We propagate this bunch in the direction of the shower in different steps. For each step, we calculate the probability to have scattering in the direction of the telescope, and the Number of scattered photons.

3)We propagate each "scattered bunch" till the telescope, Considering attenuation coefficient of the different cells the bunch pass through





