

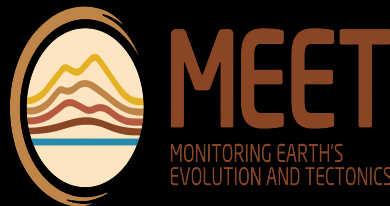


INGV

Active and passive seismic measurements at the Sardinia ET vertices

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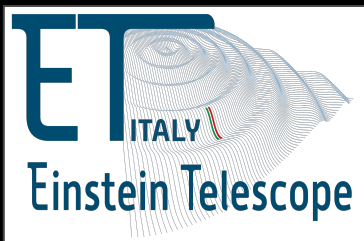
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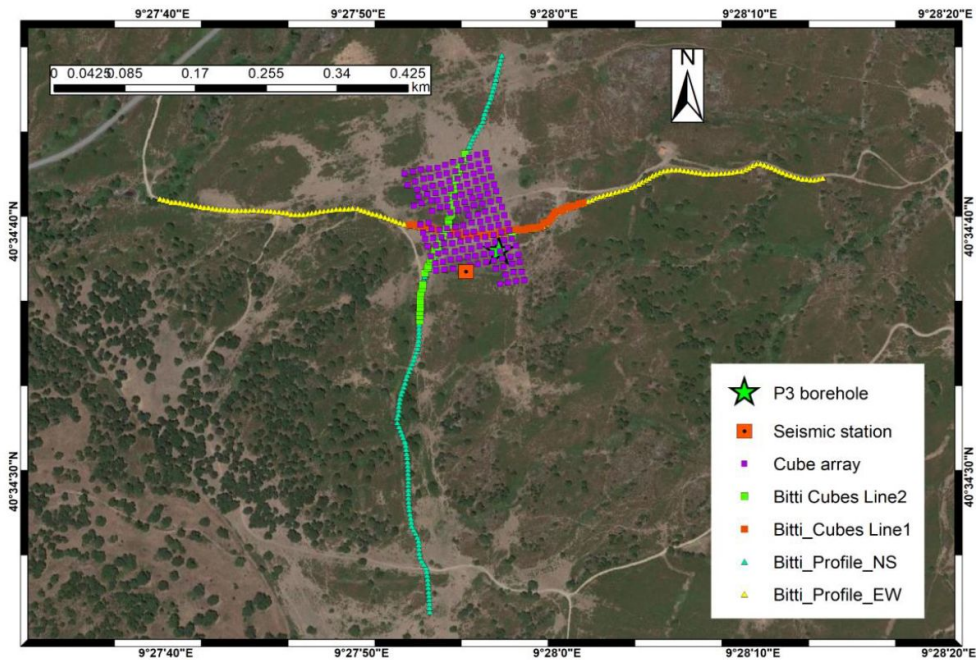


Overview of the measurements at vertices

Sos Enattos	Multiparametric network Passive Seismic array	2019 01/2021
P2	Borehole Active seismic survey Passive seismic array WINE (wind park)	09/2021 09/2021 09/2021 04/2023
P3	Active seismic survey Borehole Passive seismic array	7/2021 9/2021 10/2021



P3 CORNER



Active seismic survey@P3 CORNER



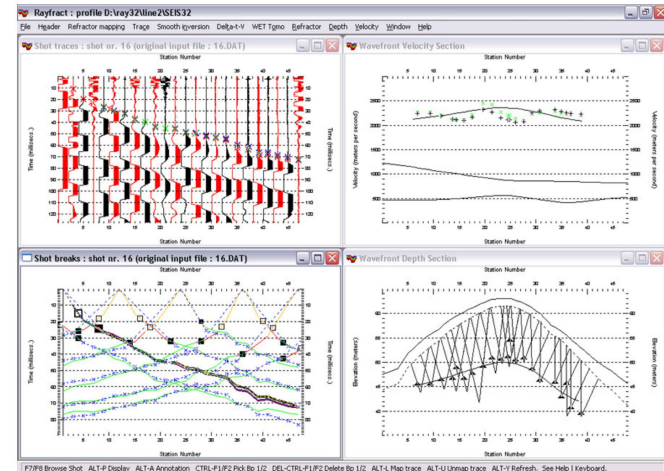
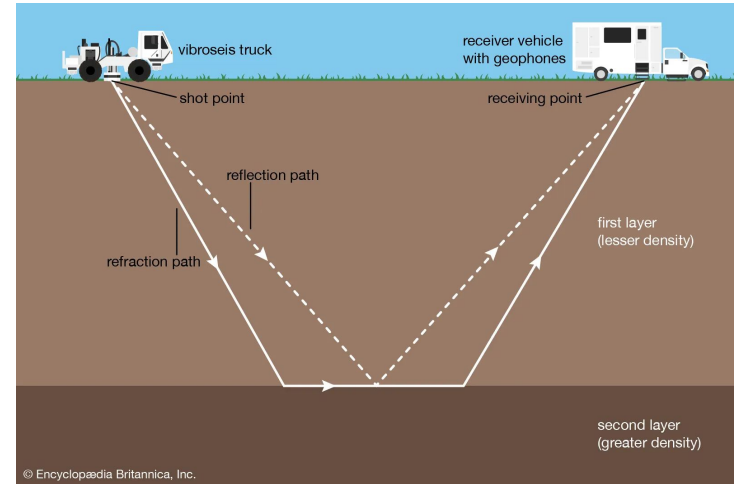
- Minivib (6 ton truck with 168 kg mass)
- up to 168 4.5 Hz geophones connected to several recorders (~ every 5 m)
- NS profile 720m
- EW profile 840m
- 1 shot every 10 m

Refraction tomography

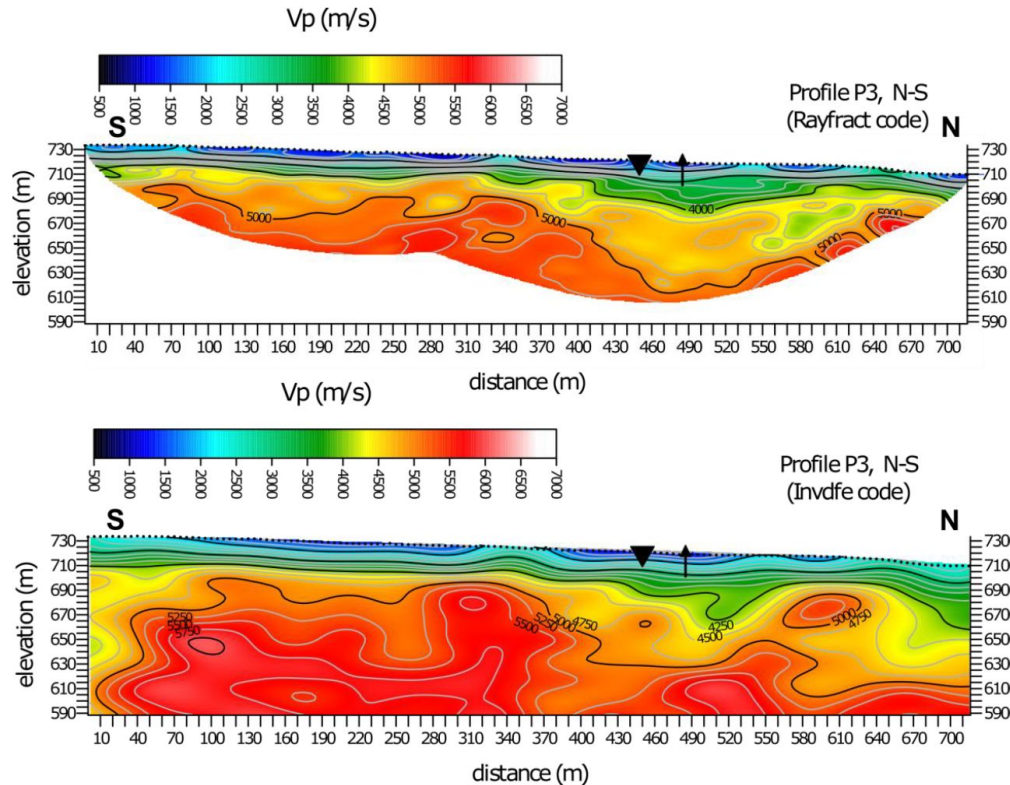
- Survey
- data gathering
- preprocessing
- pick of first arrivals
- imaging of P velocity

Inversions are performed by two methods

- Rayfract (commercial software)
 - rapid and robust V_p inversion
 - generate initial smooth velocity model
 - refinement to minimize observed delays
- INVFDE (research code)
 - multiscale approach
 - nonlinear inversion scheme (Monte Carlo + Simplex)
 - increasing number of parameters to minimize misfit
 - adaptive resolution for finer details of velocity structure

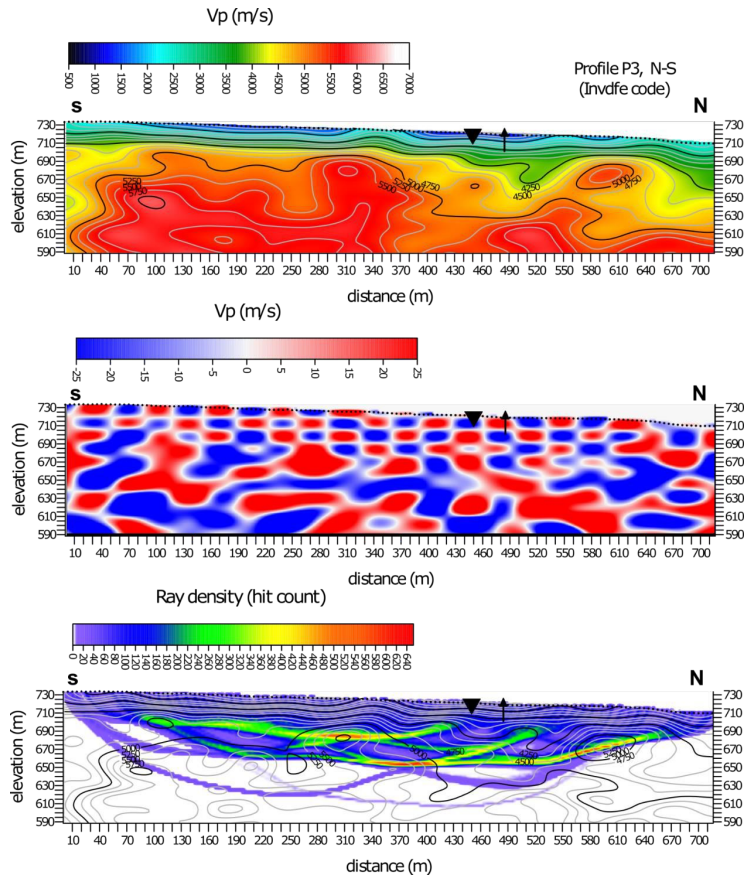


P3 NS profile: Rayfract vs Invfde (7,994 arrival times)



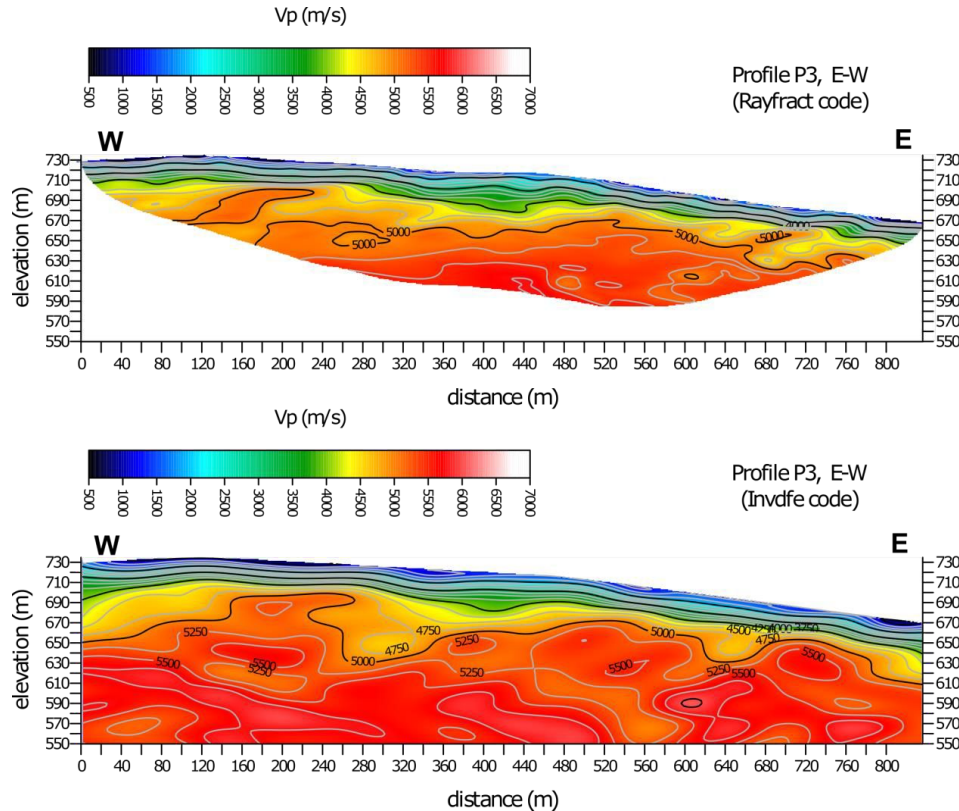
- good agreement
- better depth resolution for Invfde
- first 30-40 m are “slow”
- below 30-40 are very fast
- no significant lateral variations

Invfde resolution test



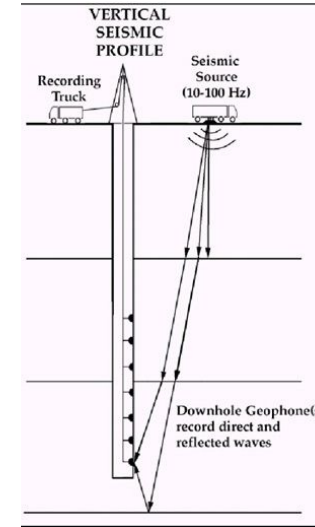
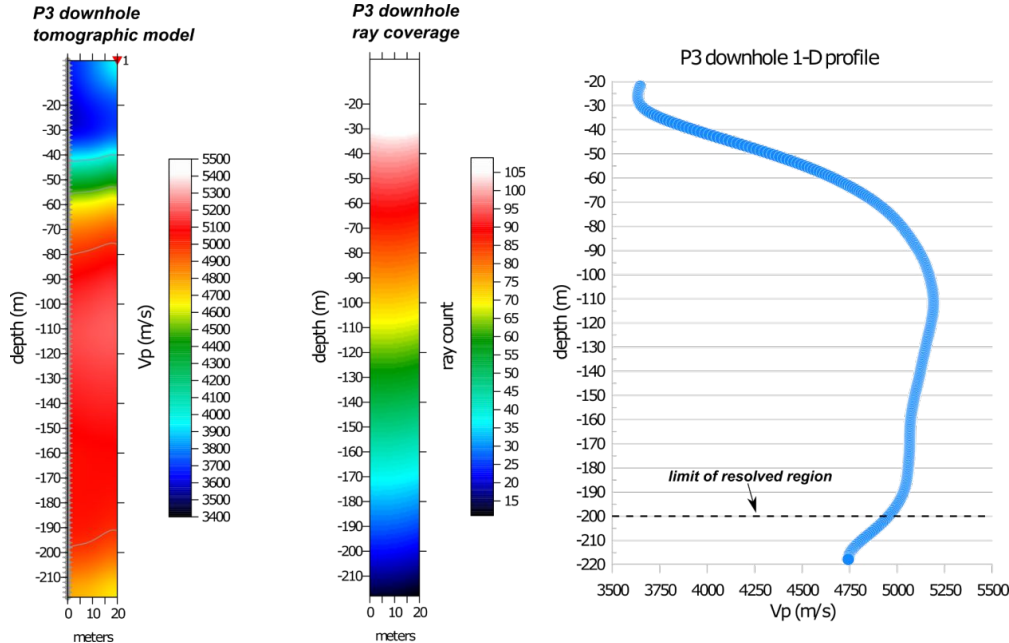
- Invfde depth resolution is validated by a checkerboard test
- Good resolution at S end and middle part of the profile (70-80 m)
- Poor resolution in the N end
- Confirmed by the ray density plot
- Fast lithologies “squeeze” rays upward

P3 EW profile (9,359 arrival times)



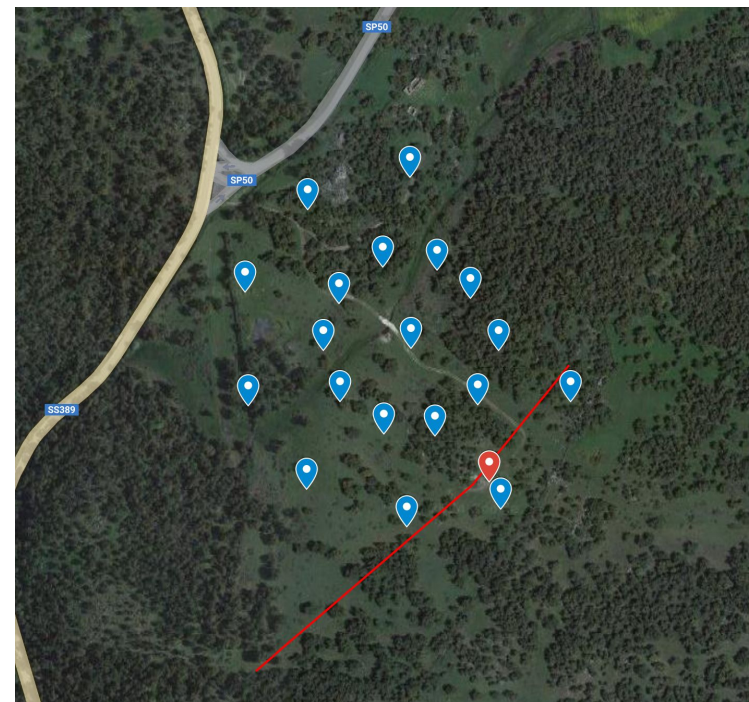
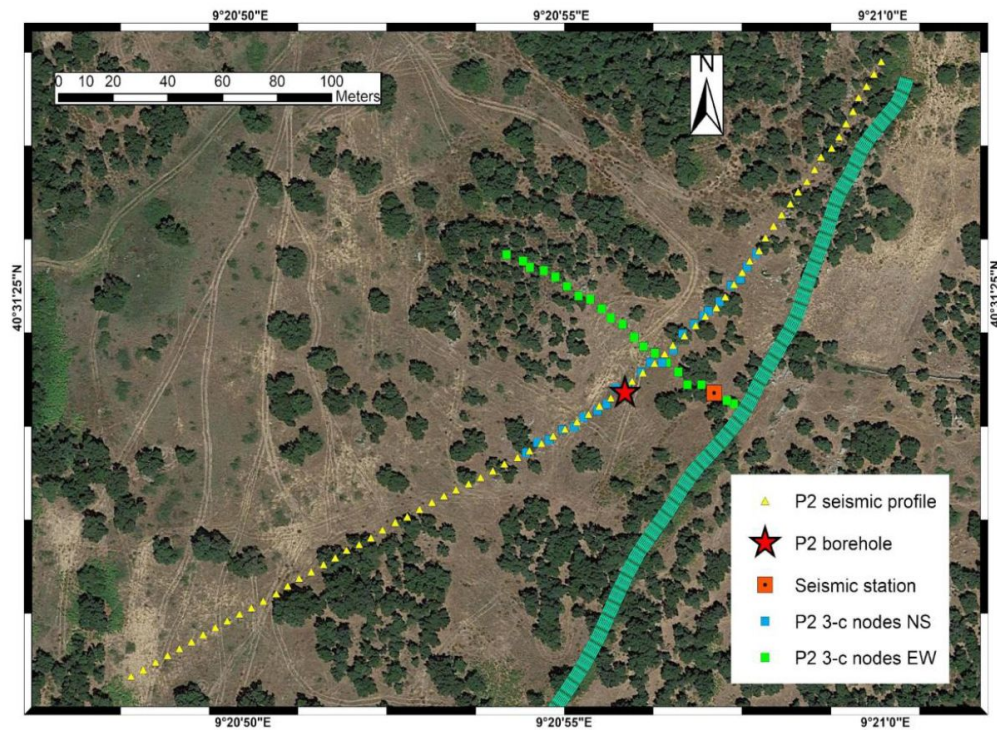
- no lateral discontinuities
- fast velocities below 30-40 m depth
- model resolved down to 50 m depth in the W part
- model resolved down to 80-90 m in central and E part

P3 corner - Downhole

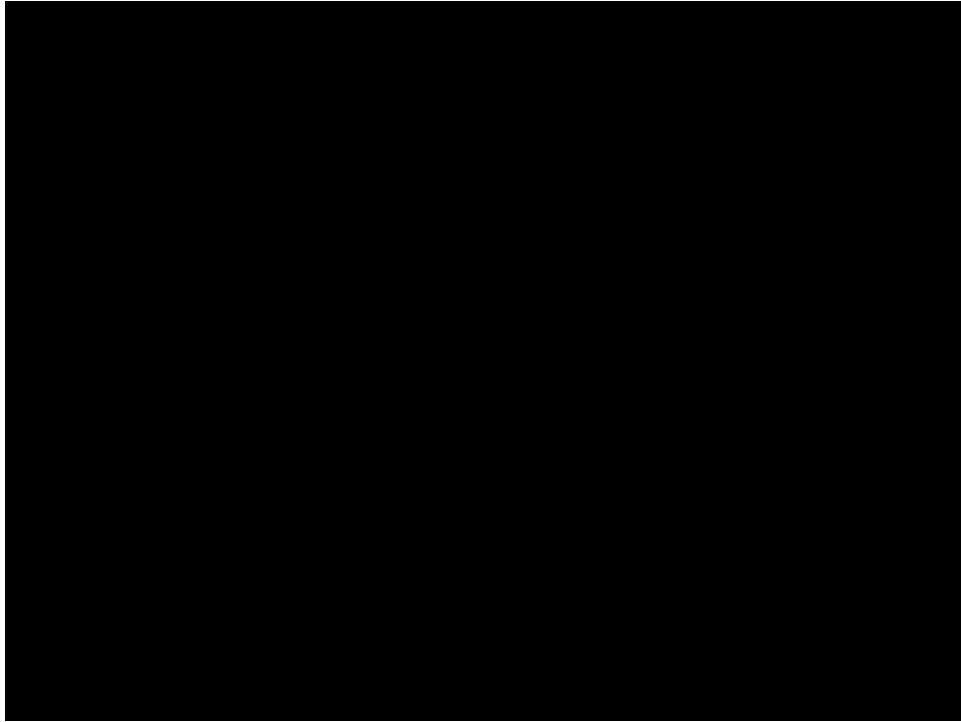


- Downhole data are processed by a tomographic technique allowing for curvature of rays
- Granitoids are altered in the first 30-40 m ($V_p < 4000$ m/s)
- Linear increase down to 110 m
- Homogeneous layers down to the resolved depth (~ 200 m)

P2 CORNER



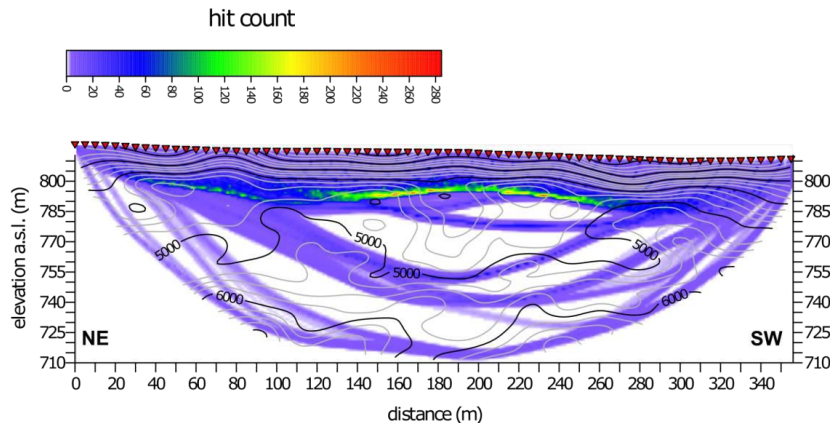
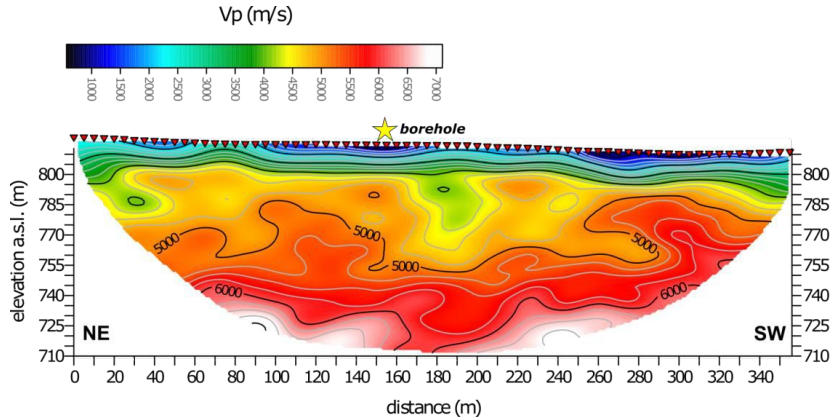
P2 CORNER Active seismic survey



- Truck unable to access the site
- Seismic gun (minibang) used as active source
- Only one profile (360 m length. SW-NE)
- 39 bangs (1 every 5 m)
- ~2,500 seismograms recorded



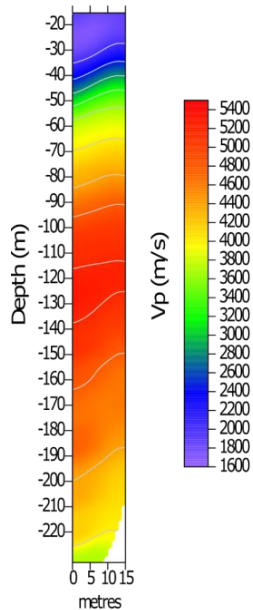
P2 SW-NE Profile



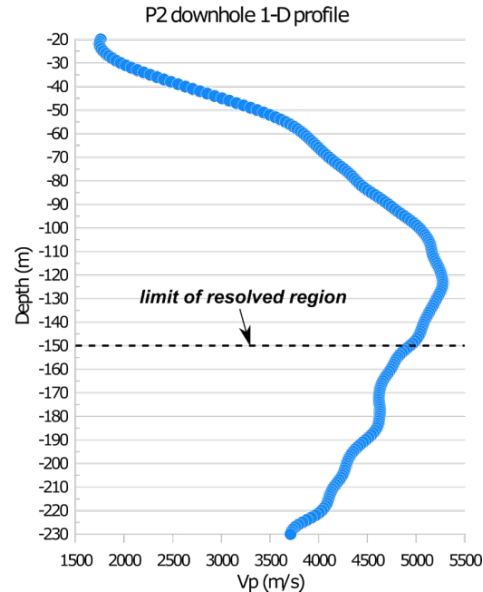
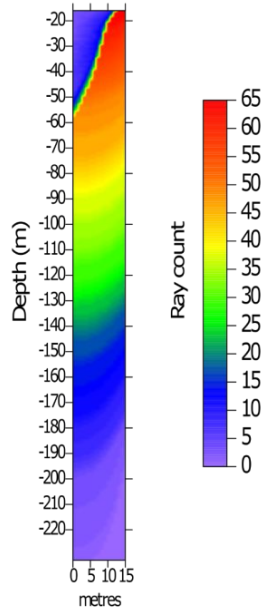
- Maximum resolution depth ~ 80m
- High velocities gradients at the surface due to alteration
- “low” velocity zone (green blob) SW of the borehole could indicate a fractured zone of about ~10 m width
- stiffer granitoids SW of the borehole (red blobs)

P2 downhole profile

P2 downhole tomographic model

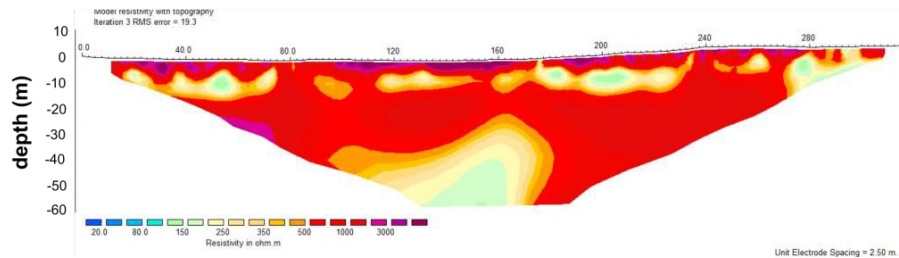
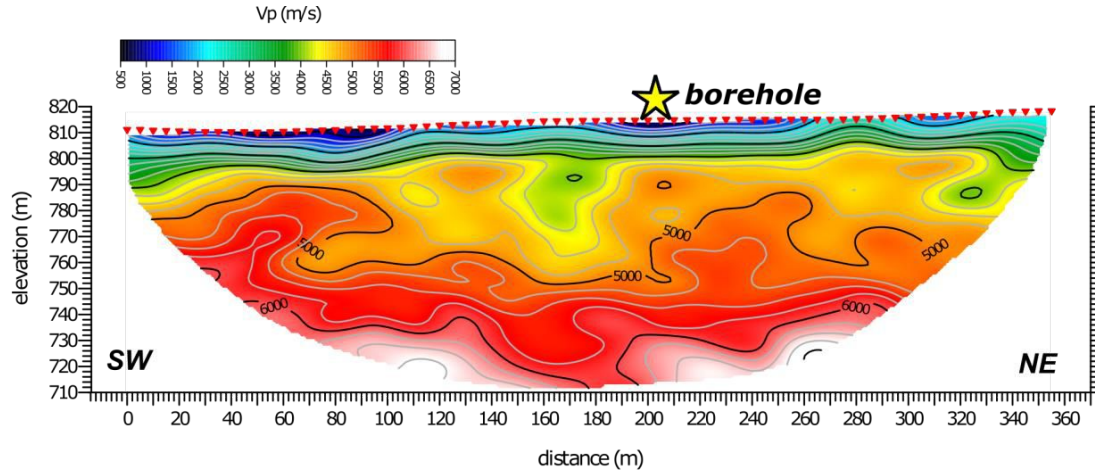


P2 downhole ray coverage



- Profile resolved to 150 m depth
- Very strong gradient to 50 m depth (2000 to 4000 m/s)
- Increase to 5500 m/s from 50 to 120m
- Slight inversion from 120 to 150 m

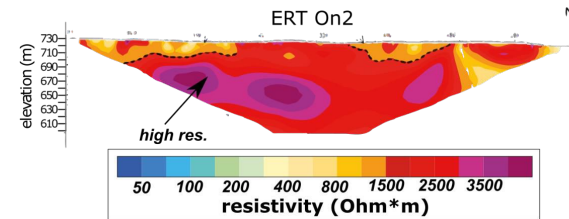
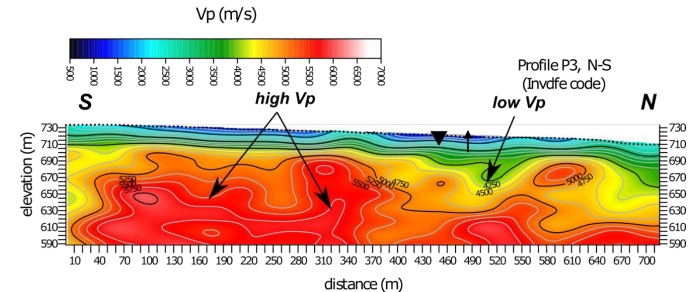
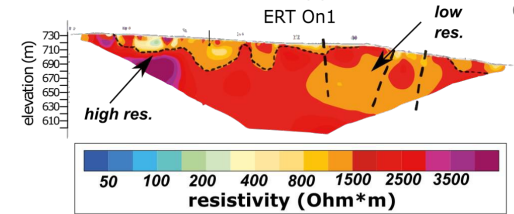
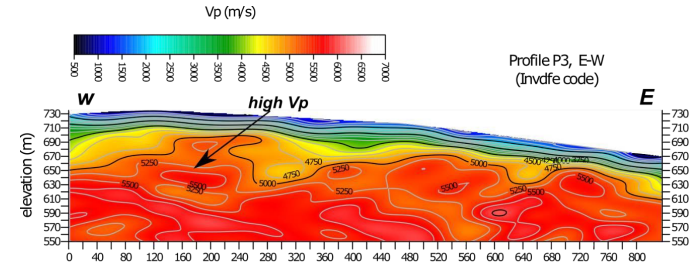
P2 Comparison with ERT survey



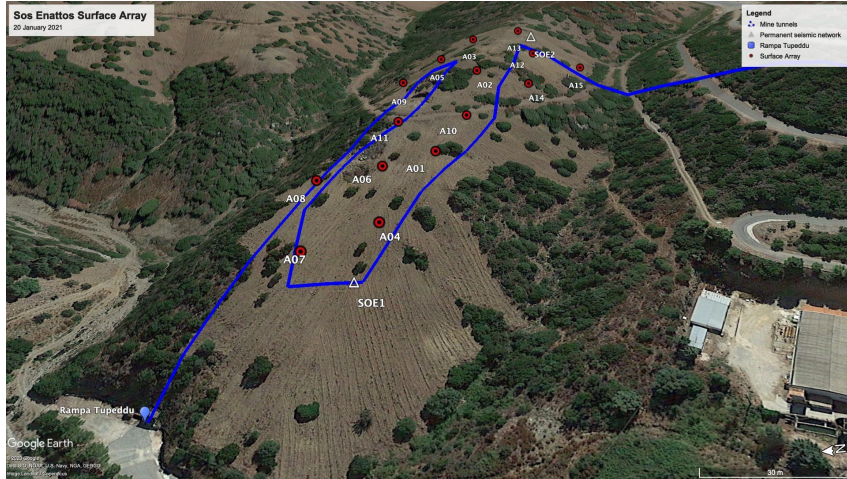
- ERT performed by UniSS (Longo & Cardello, 2021))
- 10 m altered layer (highly conductive and seismically slow)
- low fractured and stiff bedrock shallower in the SW part of the profile matches high resistivity patch
- low seismic velocity/low resistivity anomaly in the middle of the profile (confined to 50-70 m depth as shown by refraction tomography)

P3 comparison with ERT survey

- Much harder to compare since:
 - profiles not perfectly co-located
 - depth of investigation of ERT very limited at the edges of the profile
- Some general features :
 - high seismic velocity bodies match high resistivity patches
 - not all low resistivity changes corresponds to seismic velocity anomalies (e.g EW profile)
- Difficult interpretation about what is going on N of the borehole in the NS profile



Sos Enattos (P1) Passive Array



- 15 broadband seismometers
- 15 days of recording time (21/01/21 to 05/02/2021)
- array aperture 100 x 250 m
- target Rayleigh wave @ 5 Hz ($l = 400$ m)
- main target noise characterization

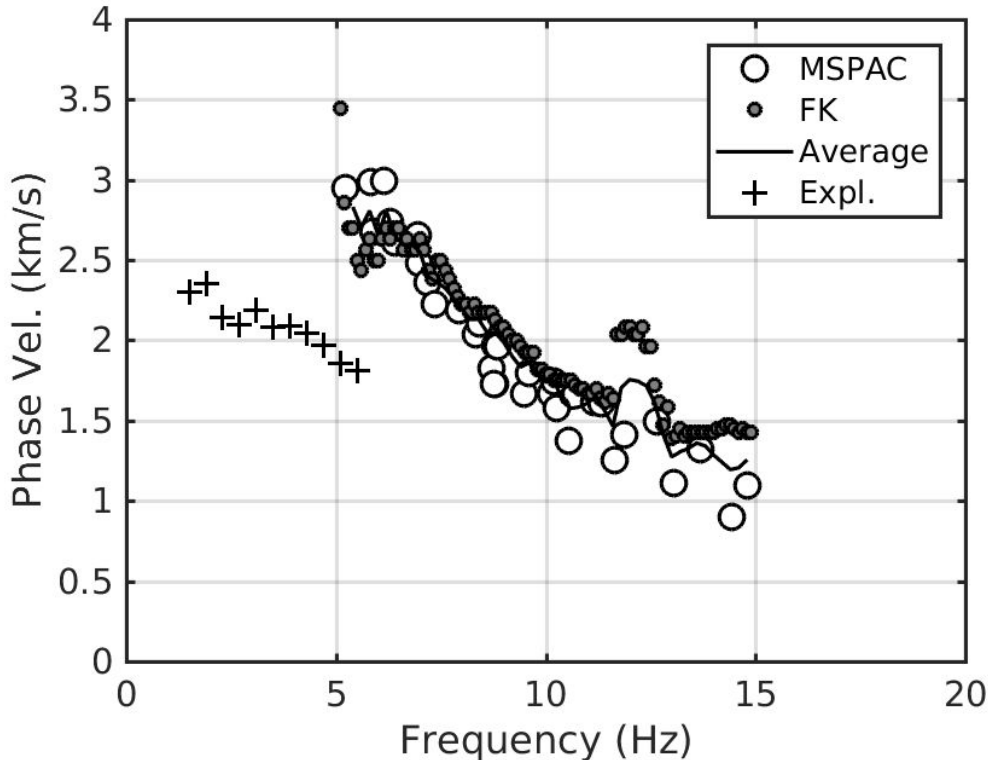
HOW IT STARTED



HOW IT ENDED

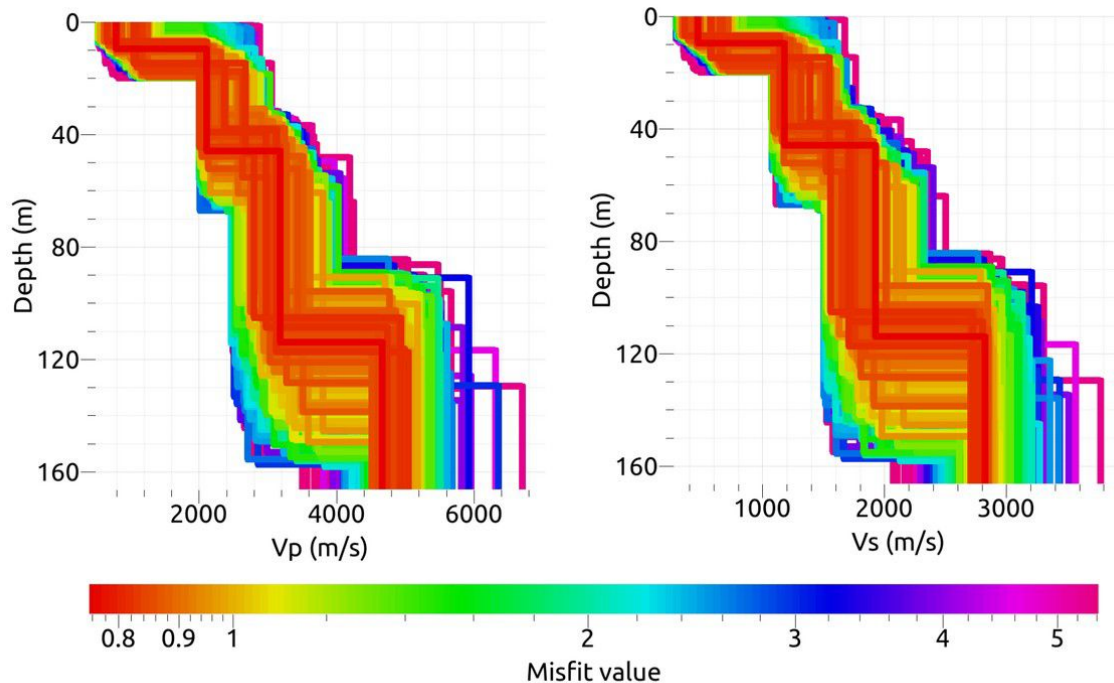


Results from ambient noise correlation: phase velocity



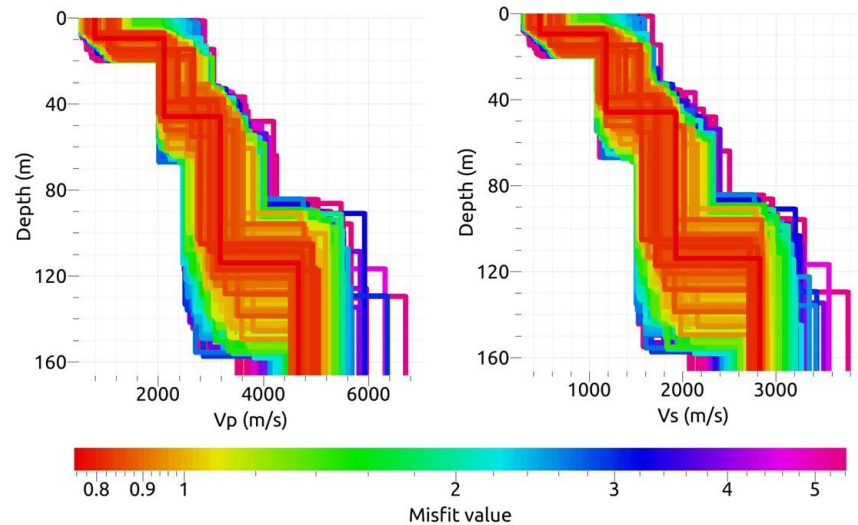
- Ambient noise interferometry: correlations of noise between stations pairs is a way to retrieve Green's functions (virtual source-receiver)
- Dispersion curve of Rayleigh phase velocity obtained by FK and MSPAC methods
- Evidence of multimodal propagation of RW

Results from ambient noise correlation: Vp and Vs profiles



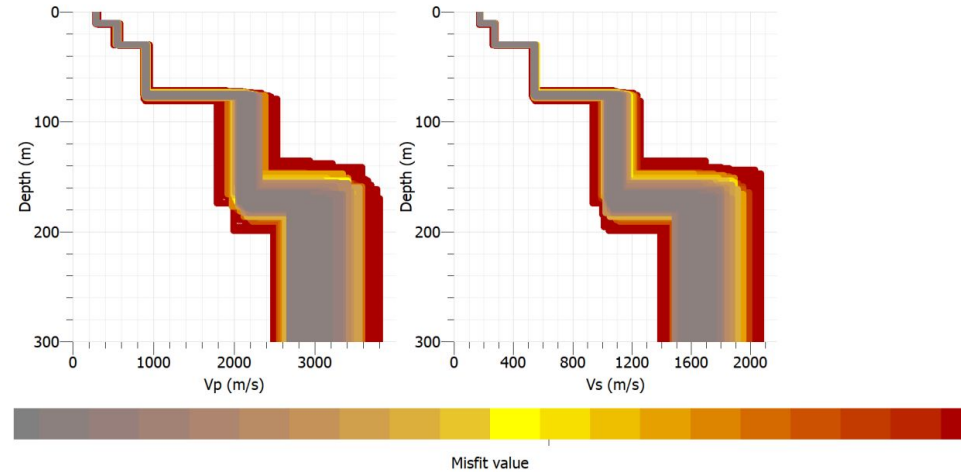
- Surface wave dispersion depends on the body waves velocities, density and thickness of the layers
- V_p and V_s inversion is performed under the hypothesis that fundamental mode and first order mode are represented by the explosion and ambient noise data
- 3 layers + half space model
- High velocities just below the surface consistent with active seismic surveys results

Comparison: Sardinia vs Euregio



Saccorotti et al., 2023

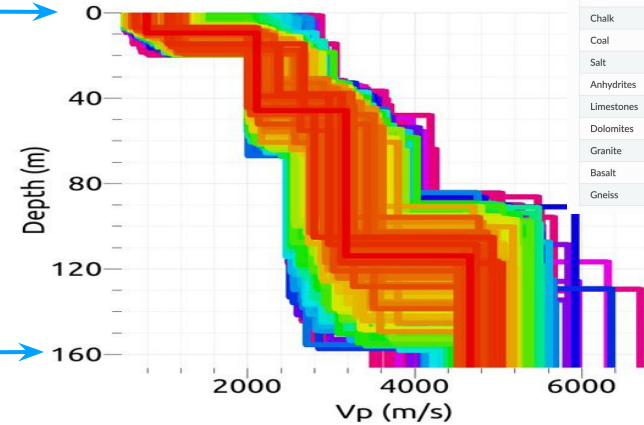
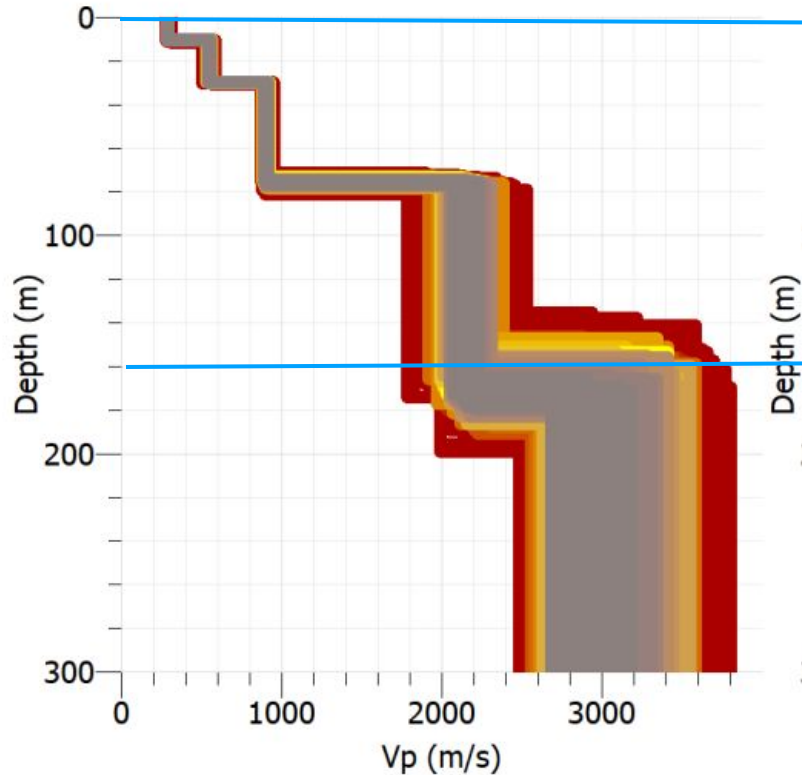
<https://doi.org/10.1140/epjp/s13360-023-04395-2>



Koley et al., 2023

XII ET Symposium, Cagliari

Comparison: Sardinia vs Euregio



Material	P-wave (m/s)	S-wave (m/s)
Air	343	N/A
Water	1450 - 1500	N/A
Ice	3400 - 3800	1700 - 1900
Oil	1200 - 1250	N/A
Vegetal Soil	300 - 700	100 - 300
Dry Sands	400 - 1200	100 - 500
Wet Sands	1500 - 2000	400 - 600
Saturated Shales and Clays	1100 - 2500	200 - 800
Porous and Saturated Sandstones	2000 - 3500	800 - 1800
Marls	2000 - 3000	750 - 1500
Chalk	2300 - 2600	1100 - 1300
Coal	2200 - 2700	1000 - 1400
Salt	4500 - 5500	2500 - 3100
Anhydrites	4000 - 5500	2200 - 3100
Limestones	3500 - 6000	2000 - 3300
Dolomites	3500 - 6500	1900 - 3600
Granite	4500 - 6000	2500 - 3300
Basalt	5000 - 6000	2800 - 2400
Gneiss	4400 - 5200	2700 - 3200

High mechanical quality of Sardinia soils is quite evident

Summary of geophysical observations at the vertices

What has been learned about the seismic velocity structure at depth?

- Both passive and active seismic studies show the presence of rather homogeneous and stiff crystalline bedrock characterized by high seismic velocities
- Very poor impedance contrasts, no major faults or discontinuities
- So far so good, but...

Towards a 3D model of the Lula/Bitti/Onanì area...

- **Doing active seismic surveys is hard in Sardinia:**
 - Logistics is complex due to topography and low anthropization (no roads)
 - Rock outcrops are very difficult to energize
 - While the depth of investigation of a refraction seismic survey is between $\frac{1}{3}$ and $\frac{1}{4}$ of the profile length in Sardinia this depth is reduced to $\frac{1}{8}$
 - Seismic surveys just scratched the surface:
 - tomography just reached 80-100 m
 - downhole reached 150 m
 - no information below 150 m (no reflections)
- **Doing passive seismic studies is equally hard:**
 - seismic ambient noise for frequency > 1 Hz is quite low
 - intrinsic noise of geophones maybe higher than ambient noise level
 - doing ambient noise interferometry with broadband sensors is a huge effort
 - same logistics problems as for active seismic survey

Airborne surveys (gravimetry and EM) could help!

What's next: actions, perspectives, and goals

- Adriarray experiment will provide some constraints on the lithospheric structure of Sardinia
- Passive seismic arrays data at P2, P3 have still to be processed (G. Diaferia): we lack dedicated full time geophysicists
- Geophysical/geotechnical investigations carried out in ETIC WP6 will focus on the design infrastructure depth (250 m)
- A full 3D model of the Lula/Bitti/Onanì area requires probably additional investigations
- Broad scale characterization of Sardinia upper crust is an interesting scientific topic for geology and geophysics