

Seismic and Newtonian noise in underground GW detectors

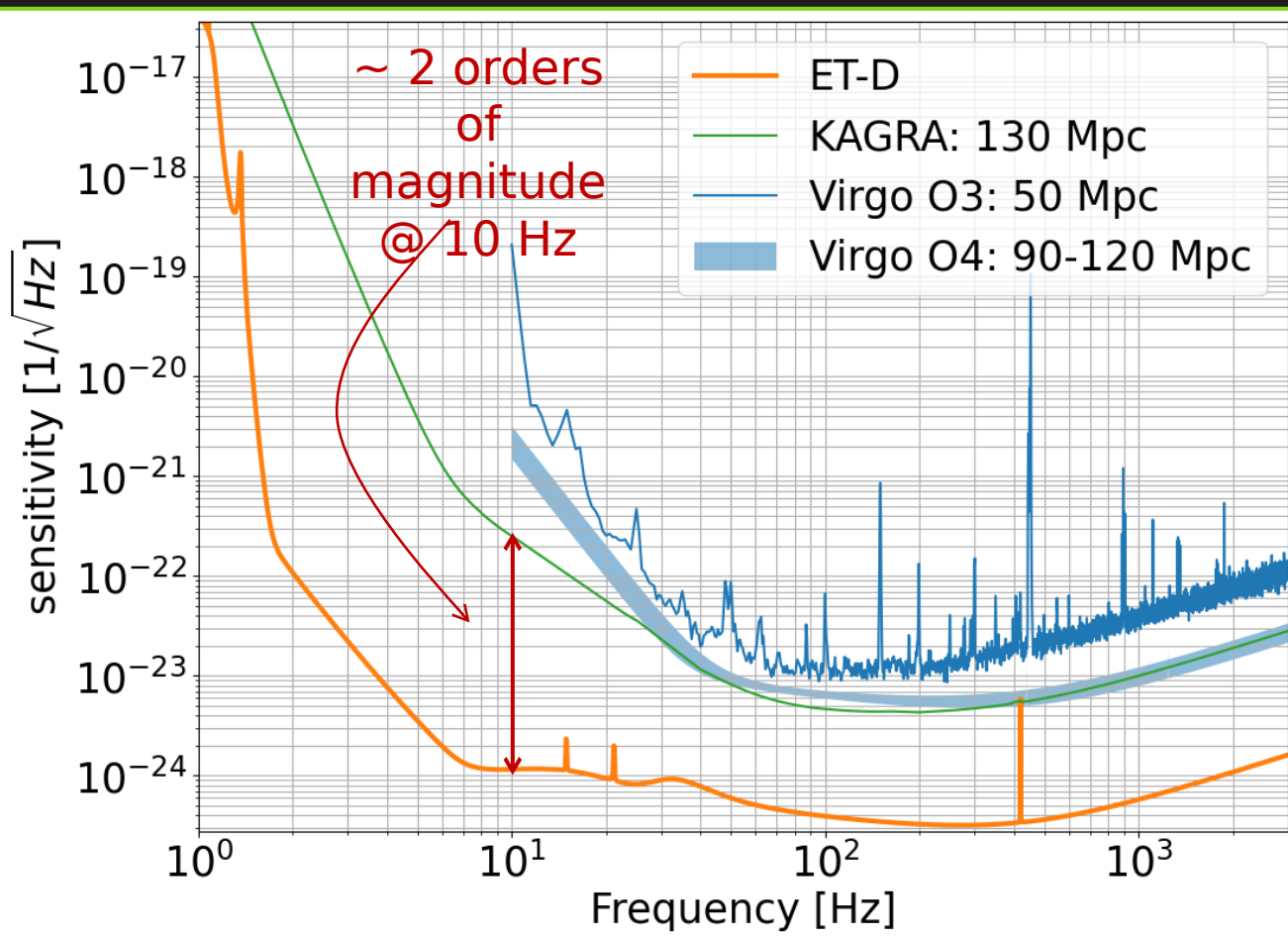
ET-Site Studies and
Characterization, November 2021

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Main goals:

- Artificially induced seismic noise propagation
- Newtonian noise estimation

Improving the **low frequency** band is very expensive: do we really **need** it?



New possible discoveries

BNS: Hours – Days
Parameter estimation
EM early warning
Sky localization with only ET

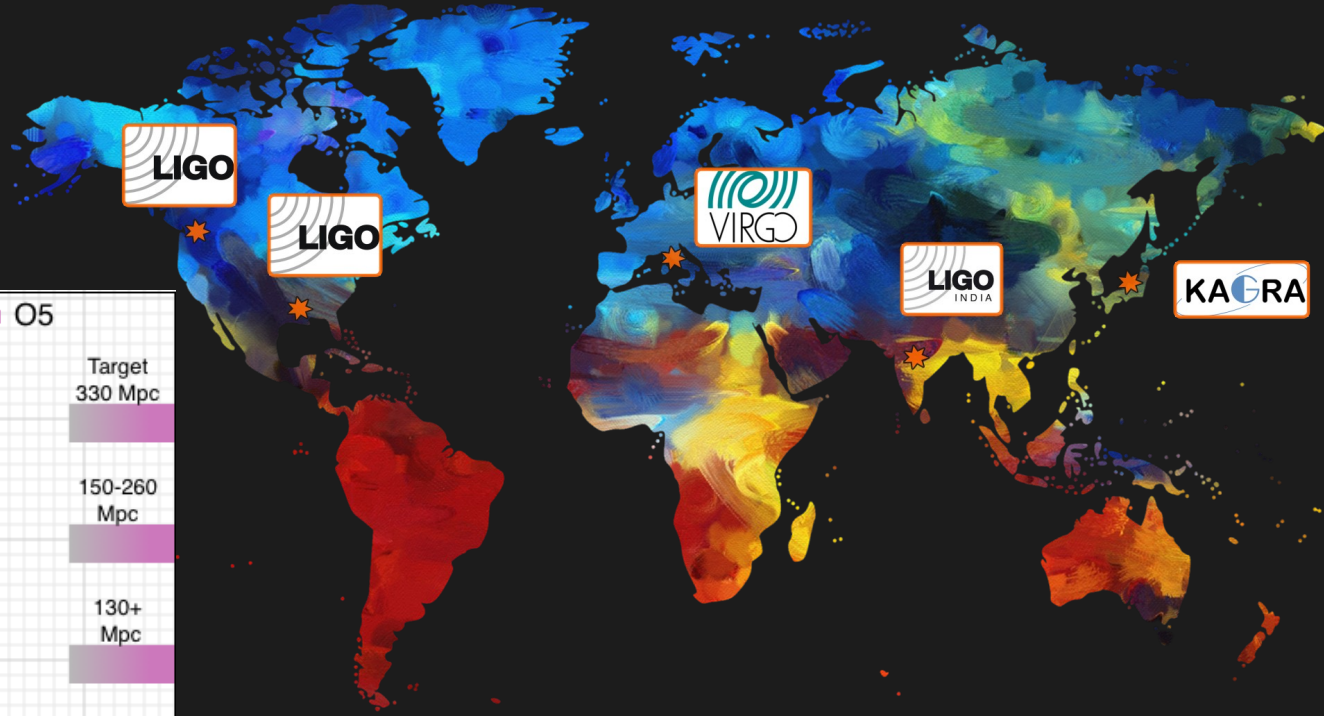
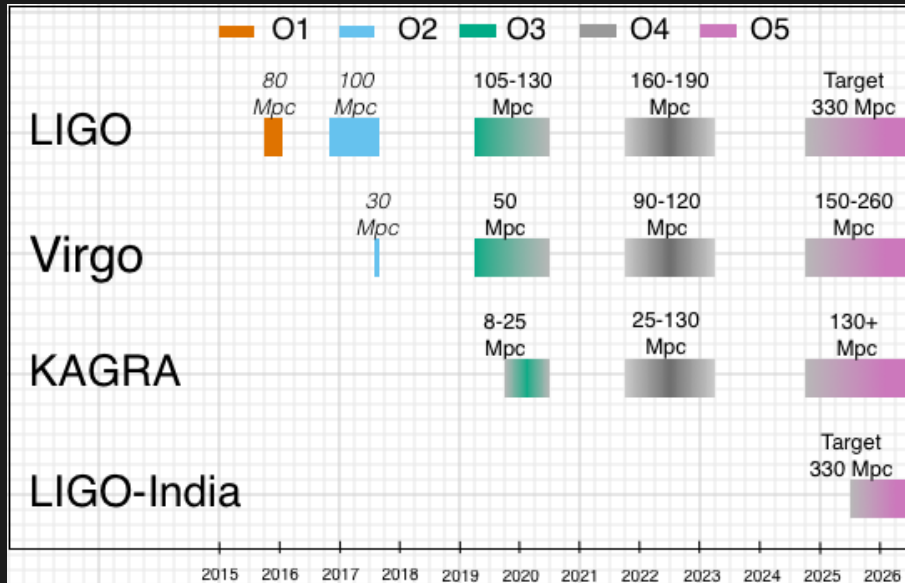
Massive BBHs:
Higher redshift PBHs?

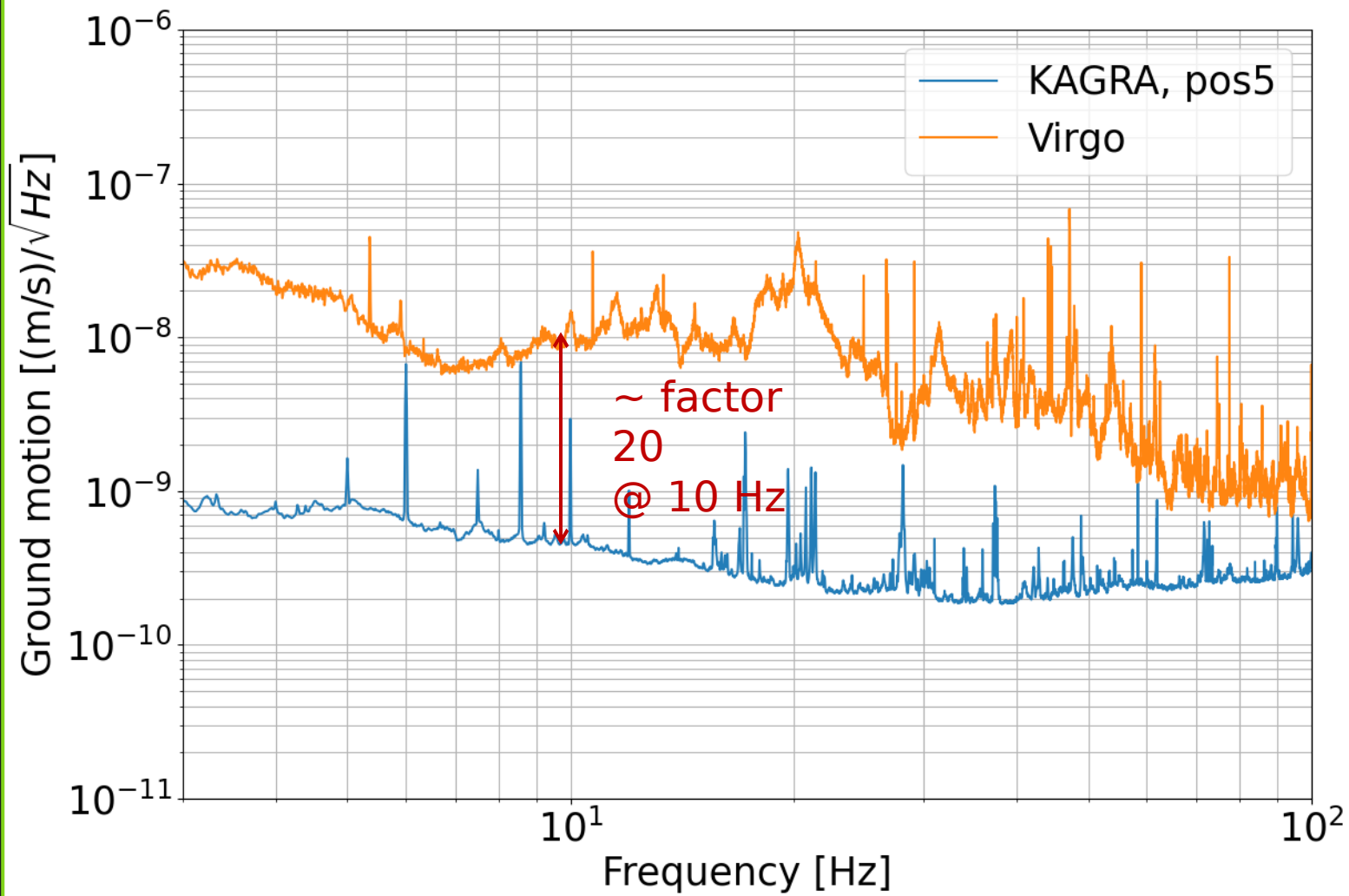
Search of stochastic background

More stable interferometer!

KAGRA is part of a worldwide
network of gravitational-
wave detectors

But it has
something special...



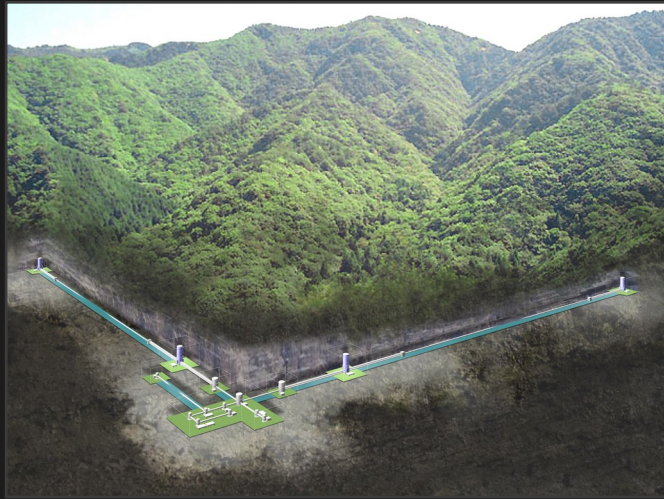


It is built
underground

...KAGRA plays
an **important**
role in the
R&D studies
for the
technologies
that **ET** will
need

And we want to
study it

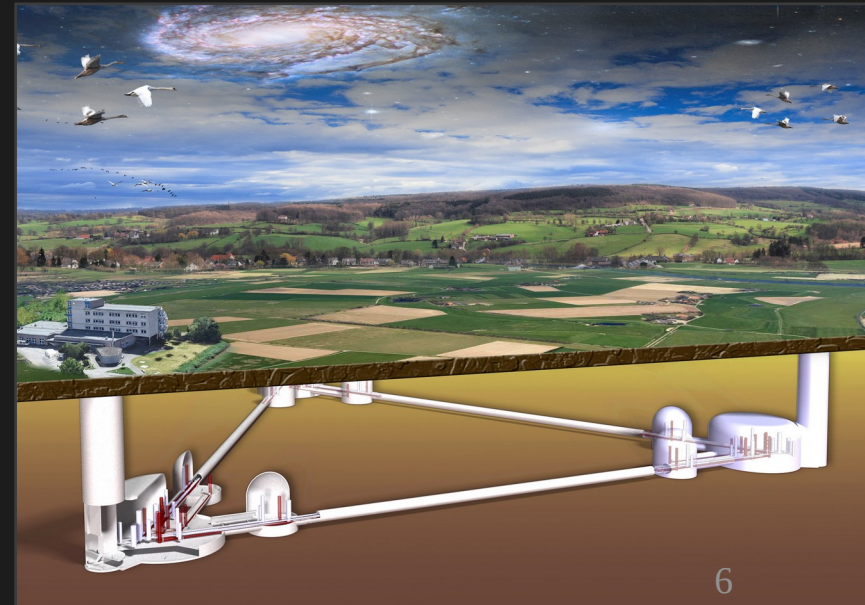
Especially
from the
Newtonian
noise point of view

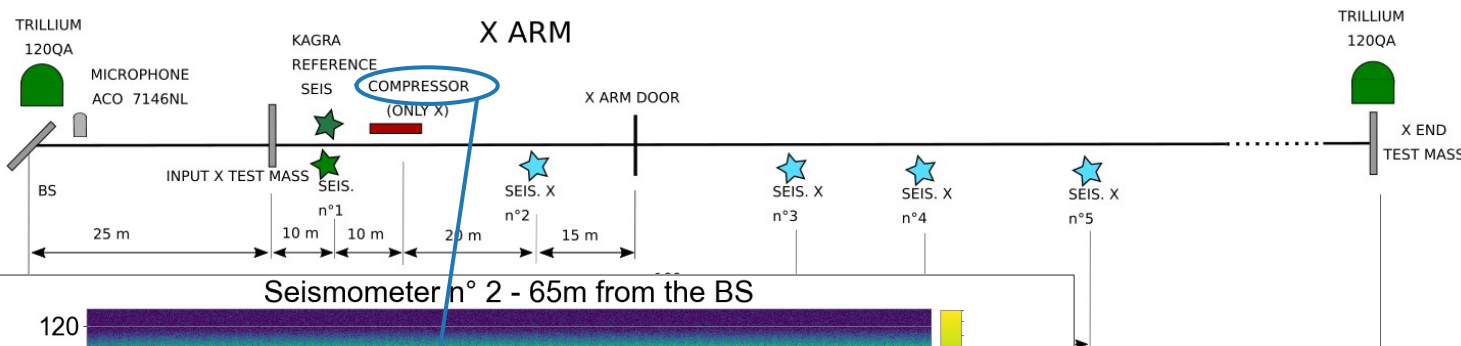


Questions:

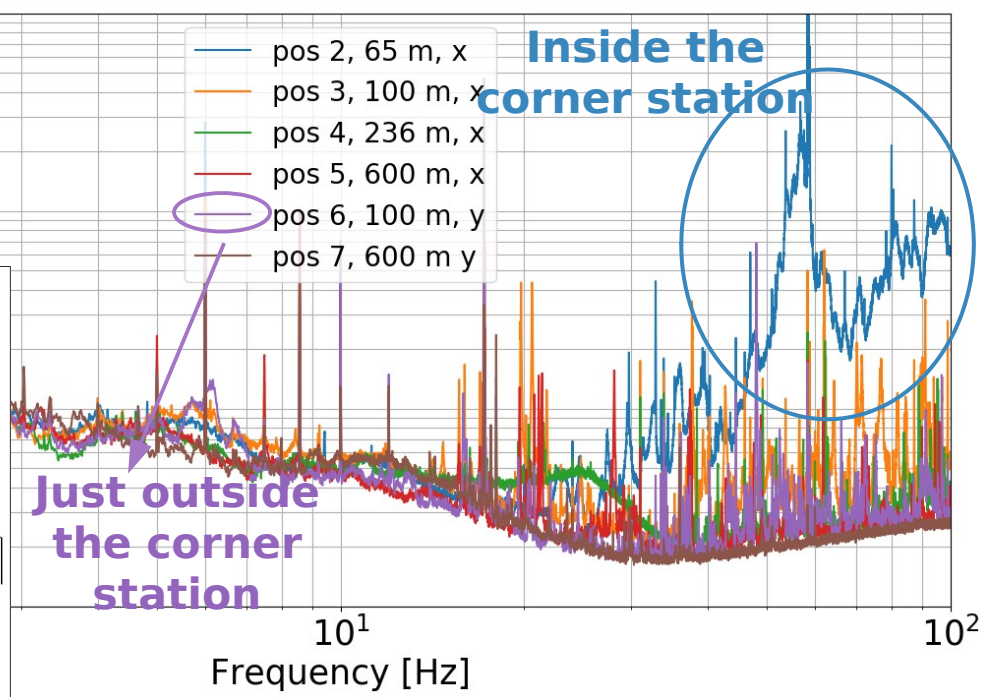
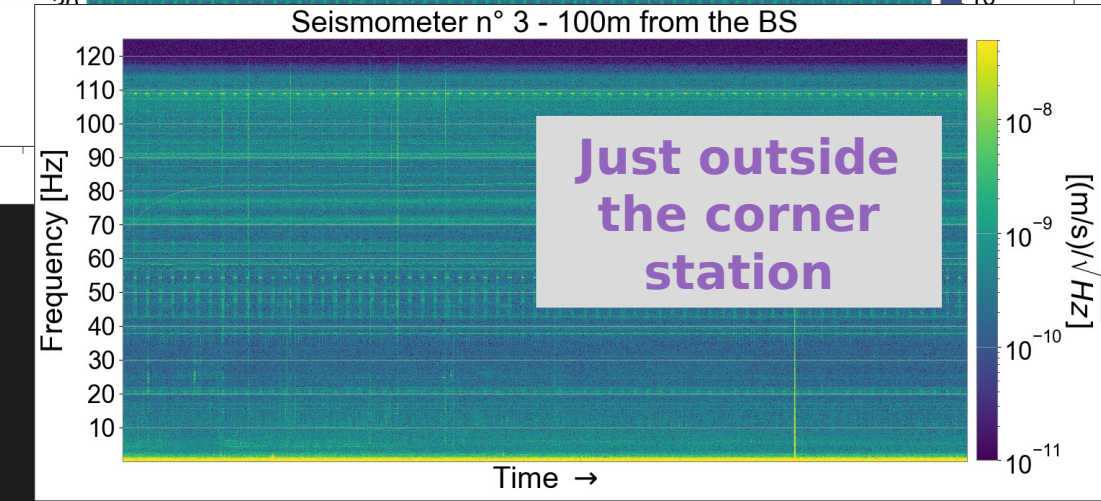
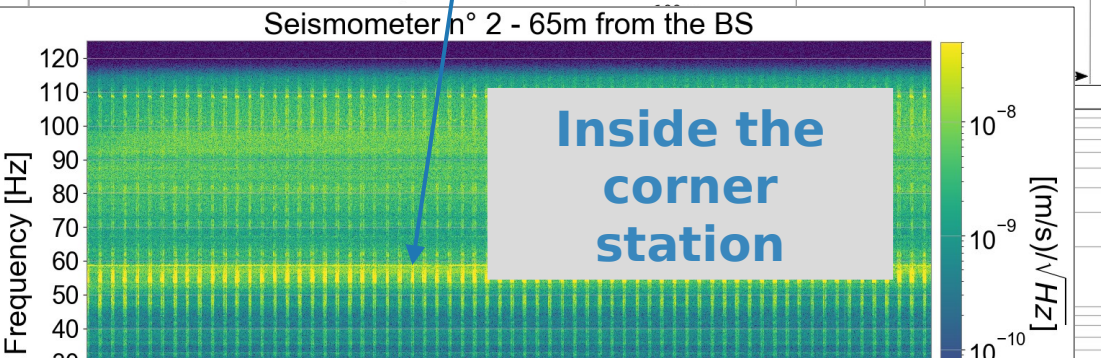
- 1) How much noise does the **underground infrastructure** cause?
- 2) How high the **Newtonian noise budget** is?

Answers:





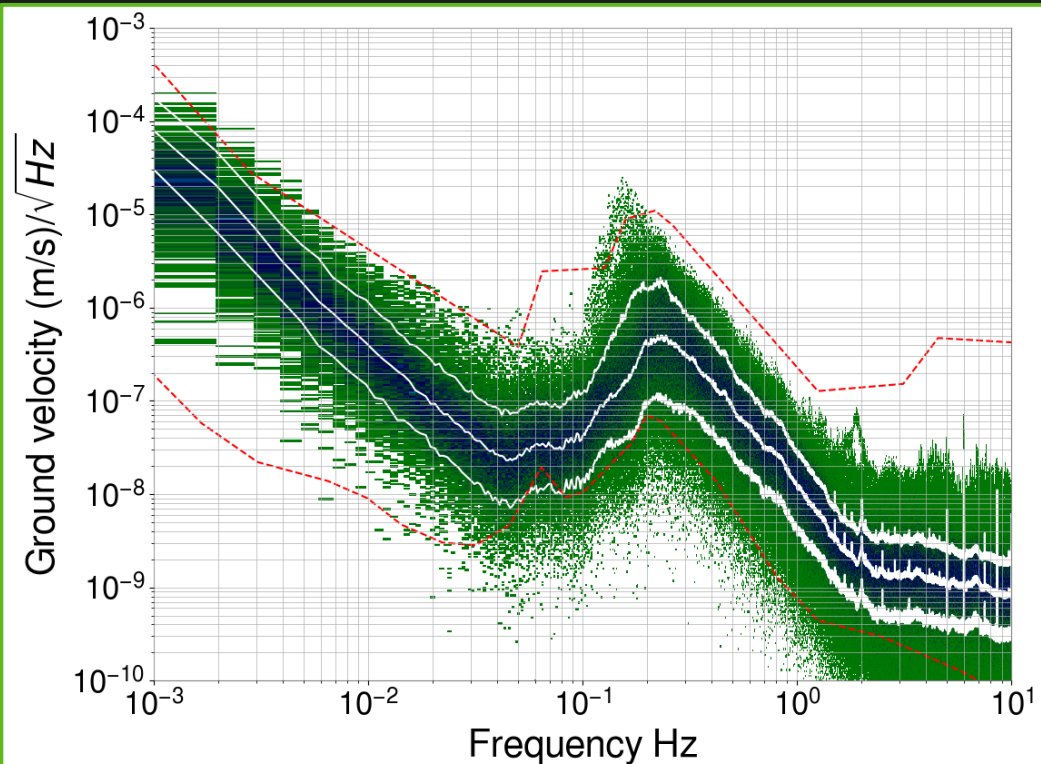
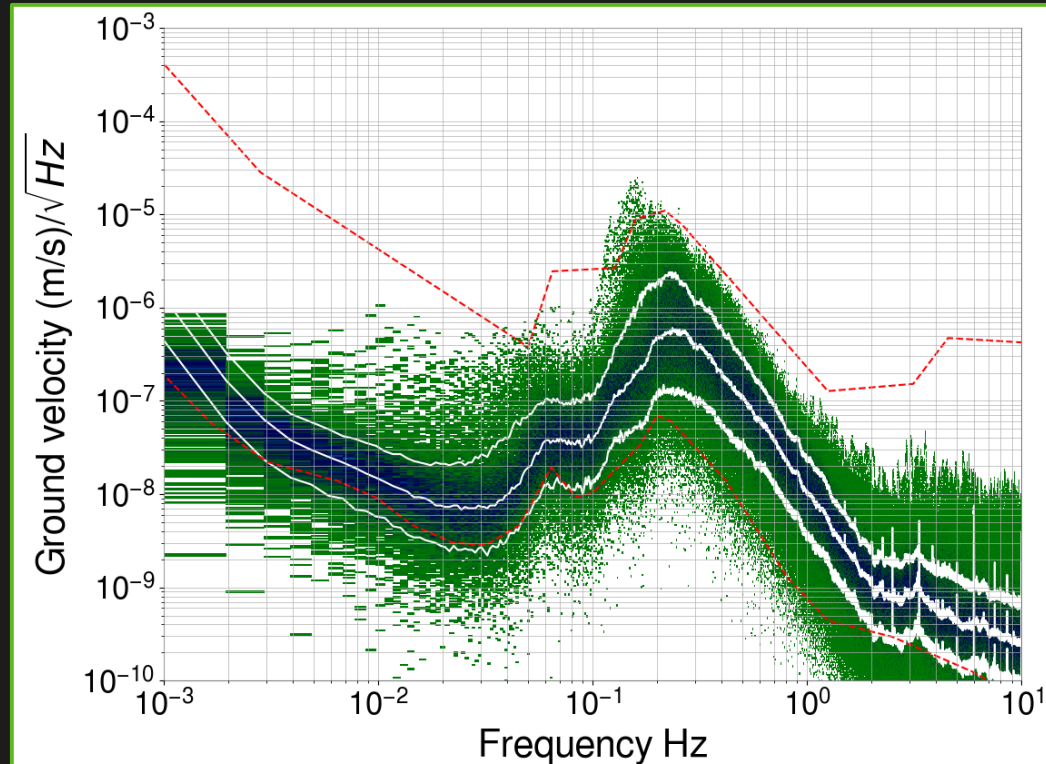
Seismic noise induced by all the infrastructures:



To be inspected: presence of excess horizontal displacement below 50 mHz.

Long-known fact: some pressure fluctuations (or thermal induced effects) Can produce seismic tilt that can couple with horizontal seismic channels.

Problem: tilt might cause troubles to the active seismic isolations of interferometric GW detectors.



$$\delta\rho_{\text{temp}}(\mathbf{r}, t) = -\frac{\rho_0}{T_0}\delta T(\mathbf{r}, t)$$

$$\delta\rho_{\text{press}}(\mathbf{r}, t) = \frac{\rho_0}{\gamma p_0}\delta p(\mathbf{r}, t)$$

ATMOSPHERIC NN

Adiabatic
index

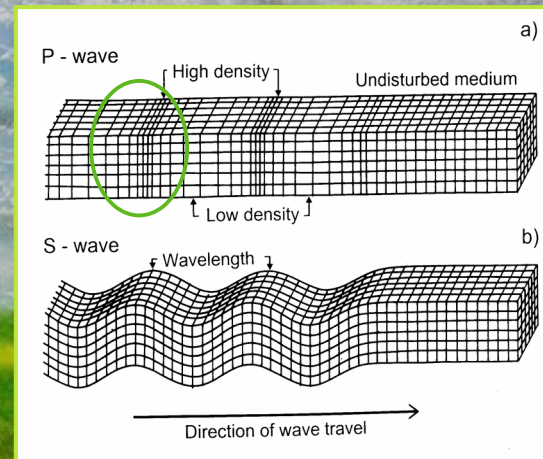
$$\delta\phi(\mathbf{r}_0, t) = -G \int dV \frac{\delta\rho(\mathbf{r}, t)}{|\mathbf{r} - \mathbf{r}_0|}$$

Newtonian Noise (NN):
Perturbation of the gravity field due to a variation in the density ($\delta\rho$) of the surrounding media.

$$\delta\rho \rightarrow NN$$

$$\delta\rho_{\text{seis}}(\mathbf{r}, t) = -\nabla \cdot (\rho(\mathbf{r})\xi(\mathbf{r}, t))$$

SEISMIC NN



Newtonian noise
budget, 3
contributions:

1) Body waves

$$\delta \vec{a}(\vec{0}, t) = 4\pi G \rho_0 \left(2\vec{\xi}^P(\vec{0}, t) \frac{j_1(k^P a)}{(k^P a)} - \vec{\xi}^S(\vec{0}, t) \frac{j_1(k^S a)}{(k^S a)} \right)$$

$$S(\delta a_x^P; \omega) = \left(\frac{8}{3} \pi G \rho_0 \right)^2 S(\xi^P; \omega)$$

x-axis
+
only P-wave
content
+
negligible
cavern
radius $a \rightarrow 0$

2) Rayleigh waves propagating underground from the surface

$$S^R(\delta a_x; \omega) = (2\pi G \rho_0 \gamma(\nu) e^{-h\omega/v})^2 \frac{1}{2} S(\xi^R; \omega)$$

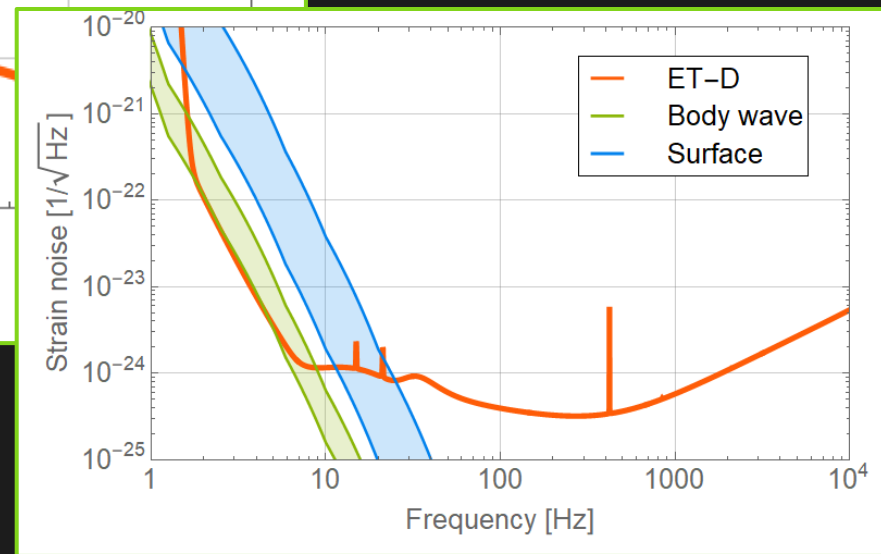
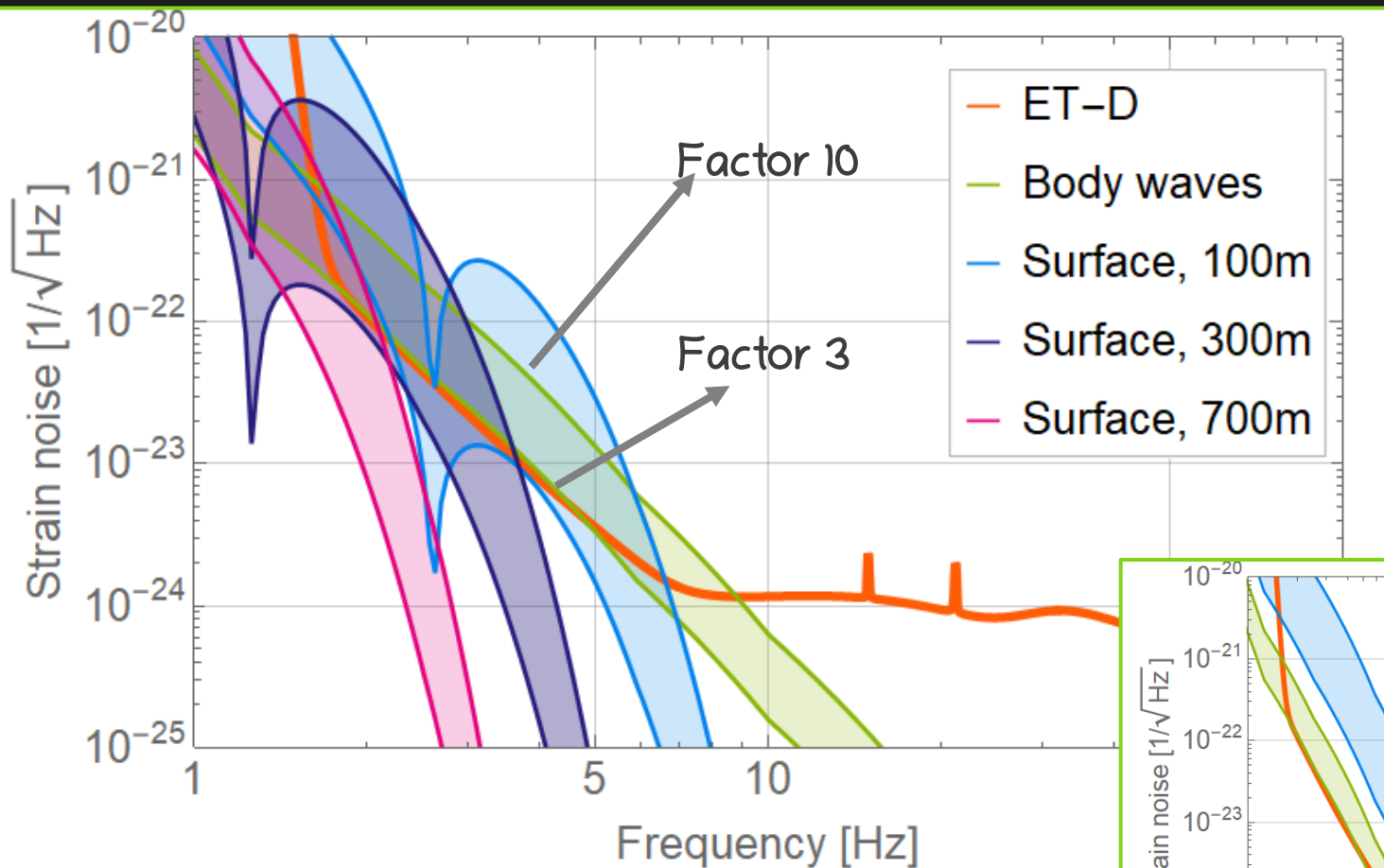
γ = determined by
the elastic
properties of the
medium

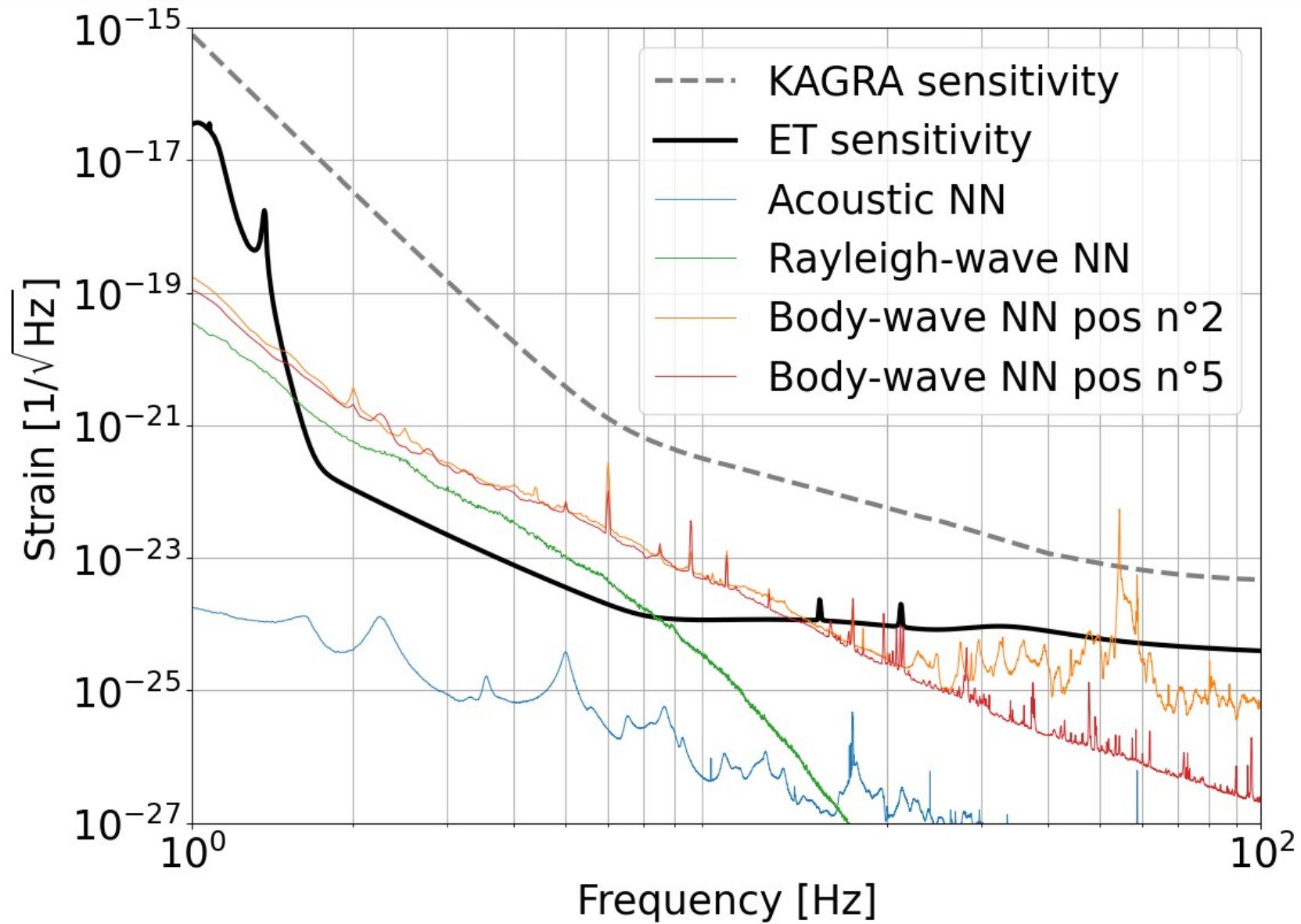
The **microseism** below 1 Hz
consists mainly of fundamental mode
of the Rayleigh waves, thus we can
retrieve Rayleigh wave velocity
from an underground array.

3) Acoustic noise

$$S_{\text{cav}}^h(f) = \left(\frac{2c_s G \rho_0 \delta p_{\text{cav}}(f)}{p_0 \gamma f} \right)^2 \frac{1}{3} (1 - \text{sinc}(2\pi f R / c_s))^2 \frac{4}{L^2 (2\pi f)^4}$$

γ = adiabatic coefficient





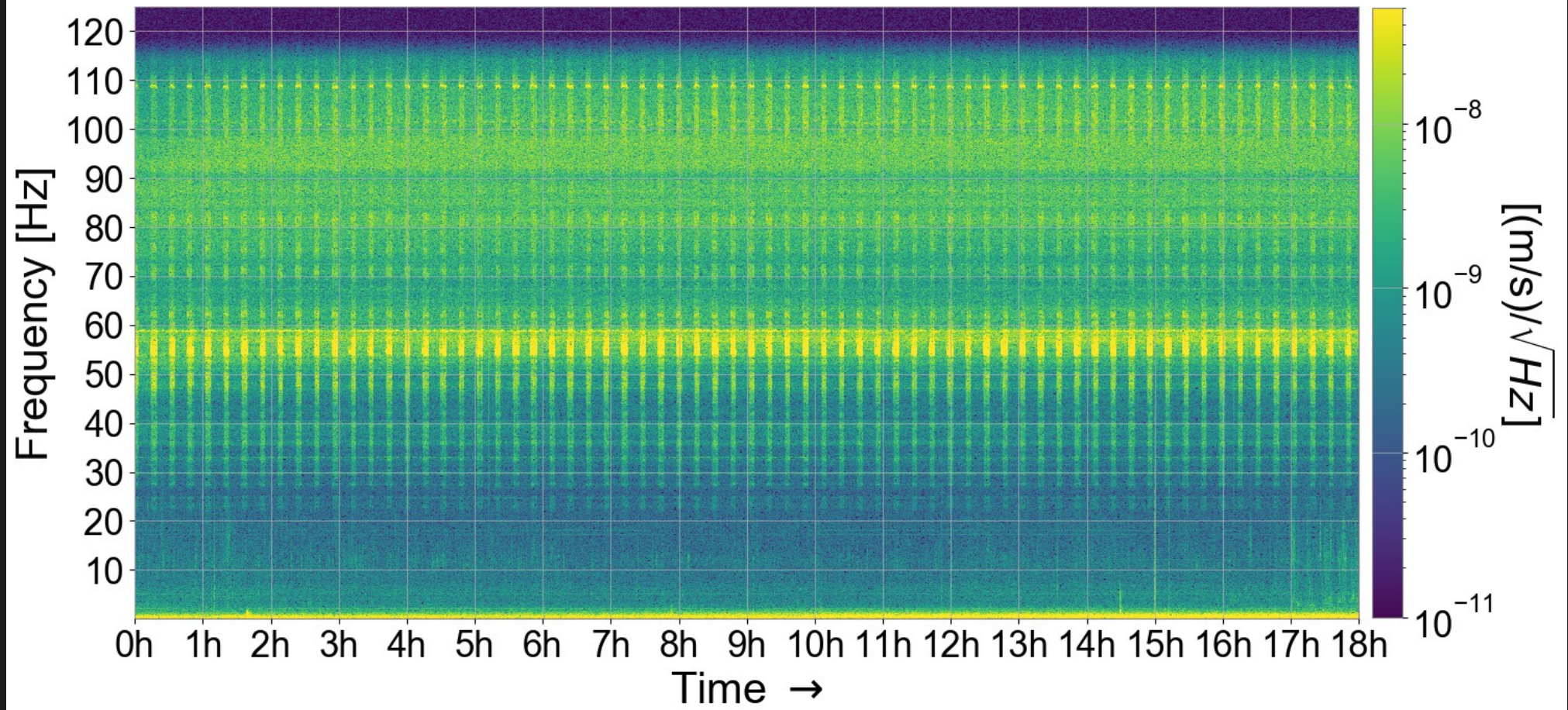
Conclusions

- Below 20 Hz, `seismic spectra` seem to be unperturbed by infrastructure sources.
- The `seismic excess` noise above 20 Hz quickly `attenuates`.
- Are the machines particularly silent or the `hard rock helps to mitigate` the coupling with the ground?
- The quick attenuation indicates that the excess noise `originates from the acoustic noise` in the corner station.

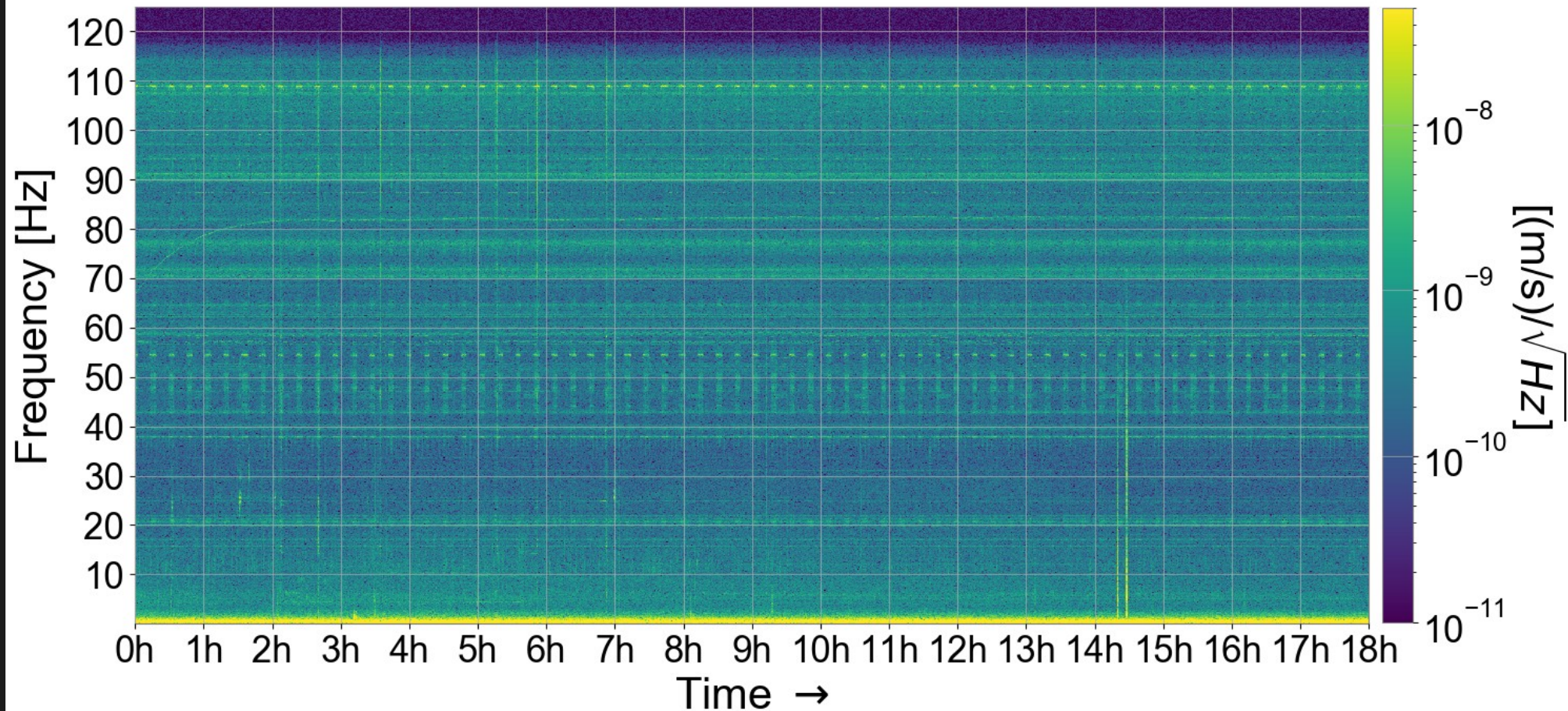
Thank you for your
attention!!!



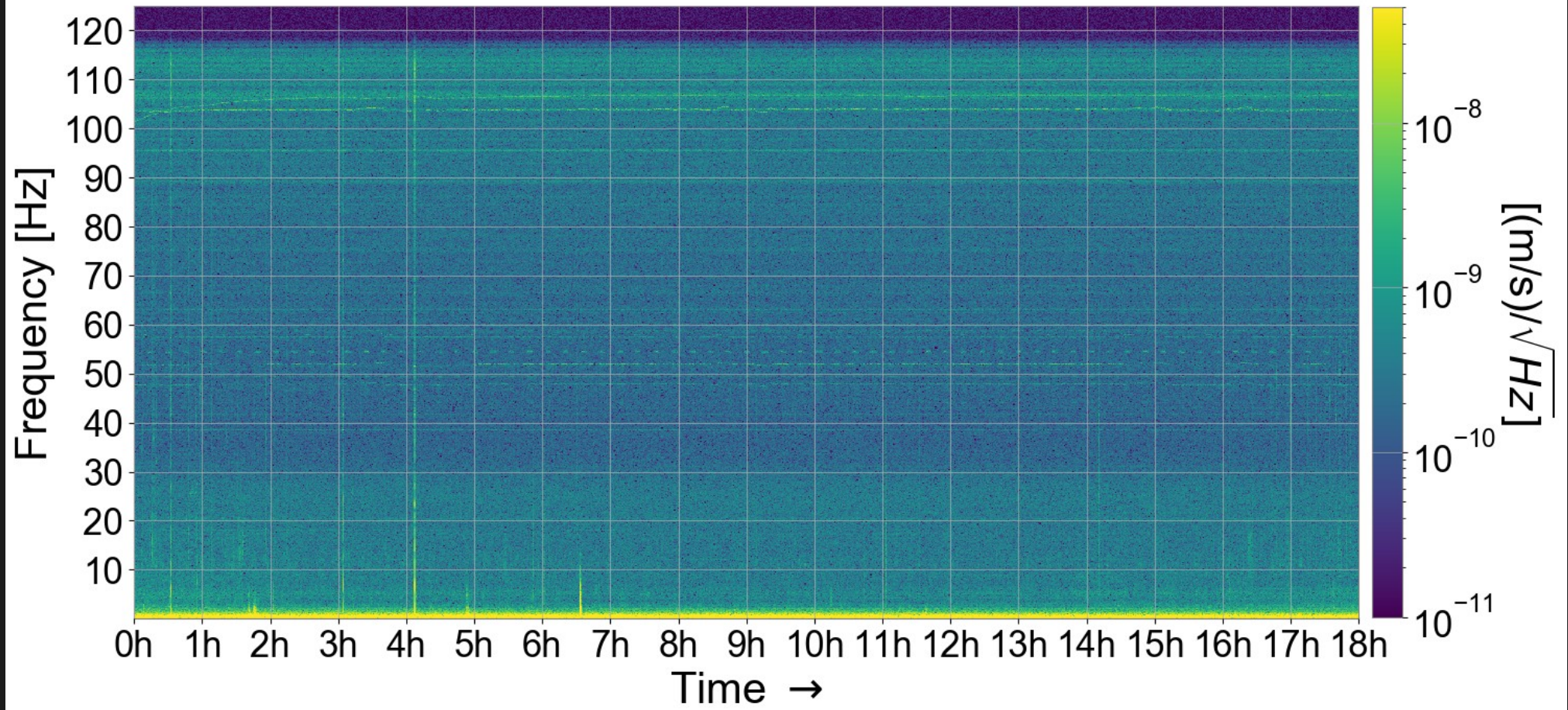
Seismometer n° 2 - 65m from the BS



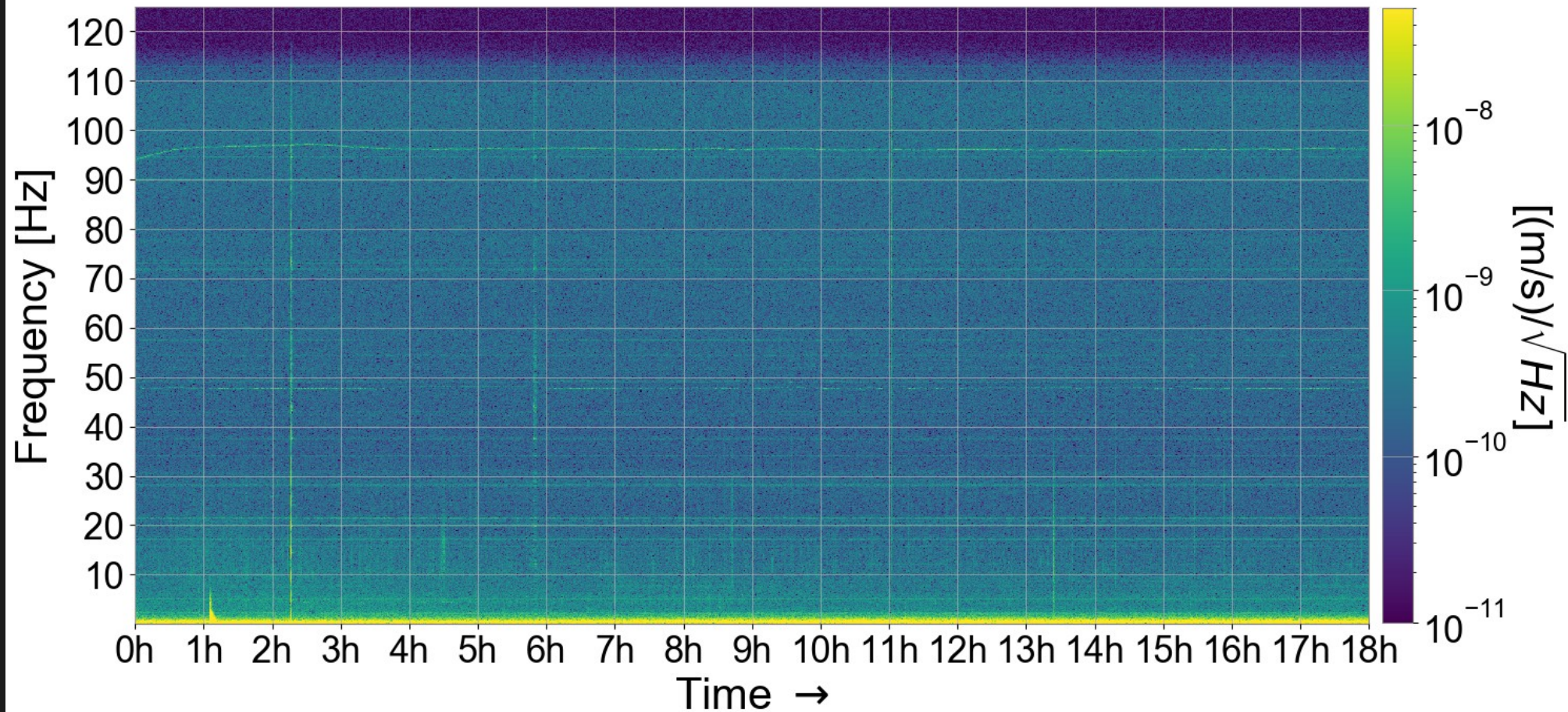
Seismometer n° 3 - 100m from the BS



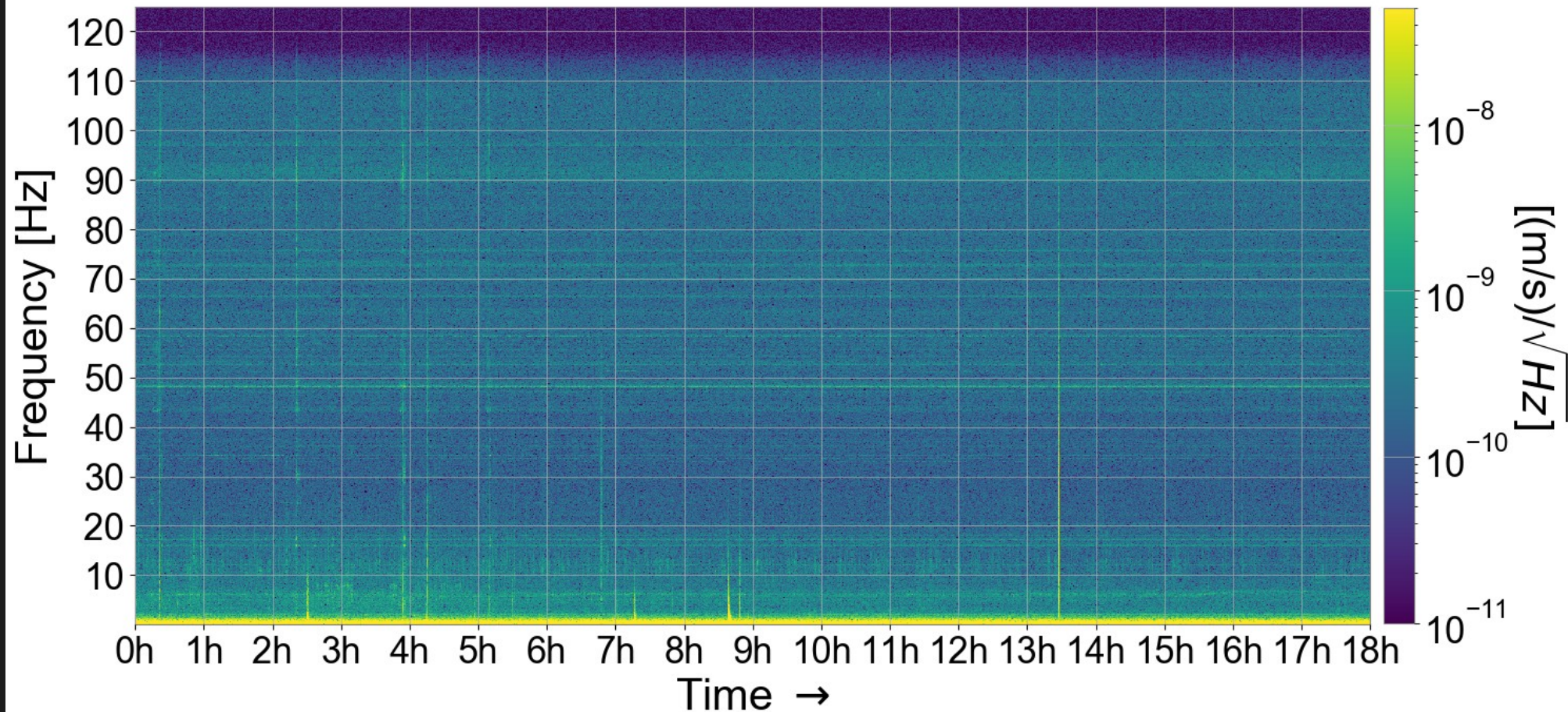
Seismometer n° 4 - 236.5m from the BS



Seismometer n° 5 - 600m from the BS



Seismometer n° 6 - 100m from the BS



Seismometer n° 7 - 600m from the BS

